

INVESTIGATING LAMINATED SEDIMENTS OF THE AL-AZRAQ BASIN, JORDAN
AND THE IMPLICATIONS FOR PALEOENVIRONMENTAL VARIABILITY

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ABSTRACT

This research examines cored sediments recovered from the Al-Azraq basin, Jordan for laminated sediments. Laminated sediments have long been recognized from the Dead Sea basin, and determined to be varved deposits, or seasonal couplets forming annual deposits. A current debate in the literature refutes the varve designation of Dead Sea sediments (Lopez et al. 2016; Ben Dor et al., 2019; Bookman, 2020; Ben Dor et al., 2020). This study focuses on the Al-Azraq sediments in a micro-scale analysis within the context of this current debate, shedding light on the potential for laminated sediments beyond the Dead Sea basin, and the implications for climate variability. The sediments analyzed from the Al-Azraq basin are in close proximity to the Dead Sea, but separated by the Jordan Plateau into an eastern drainage. It is a low lying, endorheic basin draining an area of approximately 12,700 km² of the interior desert extending from Syria to Saudi Arabia. The basin formed by the Hamza Graben is a rare deep sediment archive. Multiple environmental proxies including: lithology, smear slides, x-ray fluorescence (XRF), scanning electron microscopy (SEM), and energy dispersive X-ray spectroscopy (EDS) investigate the potential for laminated sediments and characterize the paleolake environments. Lithology and smear slides identified the presence of fine laminated sediments and diatoms. Detailed examination of ten sediment sections totaling 5.76 meters identified three sediment sections 2.02 m in length for further analyses. From these sections subsamples with high probability of laminated sediments were sent for XRF examination. XRF results of micro-scale analysis demonstrate changes in depositional

processes, alternating fine laminae of terrestrial material and Ca evaporite, possibly aragonite, and accompanying changes in moisture flux. The data demonstrate visible laminated rhythmites. XRF results identify alternating laminae of Ca and Si, SEM scans identify seasonal diatoms and EDS demonstrates an encrusting evaporite matrix. Results reveal variability in paleoenvironment with differing sedimentation above and below rhythmite sediments.

The undersigned, appointed by the Dean of the College of Arts and Sciences, have examined a thesis titled "*Investigating Laminated Sediments of the Al-Azraq Basin, Jordan and the Implications for Paleoenvironmental Variability*" presented by Jessica Hirsch, Candidate for the Master of Science Degree and certify that it is worthy acceptance.

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1. INTRODUCTION

Introduction

This study examines cored sediment samples collected from the Al-Azraq Basin, Jordan, for the presence of laminated sediments which are representative of changing paleoenvironments. These changes represent potentially significant shifts in the paleoclimate regime. This research employs methods for identifying seasonal laminated sediments used in previous research. For example at Meerfelder Maar in Eifel, Germany, X-Ray Florescence (XRF) provides micro-stratigraphical data to better illustrate the interannual, seasonal laminations of sediments (Brauer, 2008). To address the presence of laminated sediments in the Al-Azraq Basin, this study initially identified laminated sediment on the basis of the initial core description and smear slides analyses. Both methods rely on visual observation. Further analyses used both XRF and SEM scans in addition to grains size and magnetic susceptibility methods. The XRF and SEM further allowed for the exploration on a microscale of the mineral elements present in the laminated layers to further explore the hypothesis these laminations are rhythmites and of seasonal origin. This would indicate a significant shift in the paleoenvironments in this region.

Research Objective

The objective of this research is to demonstrate cored sediments within the Al-Azraq Basin record laminated sediments, and that these sediments represent several periods of cyclic deposition, potentially seasonal. A robust suite of analytical methods employed address the research objectives of demonstrating laminated sediments and cyclic deposition within the Al-Azraq basin cored sediments. The research investigates the cored sediments and elements and materials found within them through the application of: initial cored description, smear slide observation, grain size analysis, XRF, SEM and magnetic susceptibility. The research investigates species diatoms within the sediments in relation to

the potential of seasonal cyclic deposition in the basin. The research examines the clear departure from in the environments below and above the core samples. Finally, the research explores the implications of environmental changes represented by these sediments in the basin.

Research Questions

This research addresses the following questions: 1) Are there laminated sediments in the AZ2 cored sediments? 2) If laminations are present, what type of laminated sediments occur? and 3) Are the sediments reflective of environmental change? 4) How do the findings relate to regional records of similar sediments?

Research Significance

The presence of laminated sediments in the Al-Azraq Basin would be the only occurrence of laminated sediments on the Jordan Plateau and outside of the Dead Sea Basin. The next nearest occurrence of laminated sediments are in Turkey at Lake Van (Stockhecke et al., 2012). The research questions are particularly relevant to current debate in the region regarding the nature of laminated sediments (Lopez-Merino et. al, 2016; Ben Dor et al., 2019; Bookman, 2020, Ben Dor et al., 2020). The laminated sediments of the Al-Azraq basin contain information with implications for paleoenvironmental variability. These questions are also relevant to present and future regional climate by demonstrating a range of possible environments.

Thesis Organization

This thesis is organized in seven chapters. Chapter Two expands on the definition and types of laminated sediments. It discusses the occurrence of laminated sediments in arid paleolake settings similar to the Al-Azraq Basin. This chapter also addresses in detail the nature of the most relevant laminated sediments from the Dead Sea, Israel and Qarun Lake, Egypt. The chapter expounds on the current debate surrounding the new and reassessments of the Dead

Sea varves. Chapter Three describes the study area on the Jordan Plateau in terms of its geologic setting. More specifically the geologic history of the Al-Azraq basin. Chapter Three also discusses the present climate conditions and sediment history. It presents a review of sediment chronology from the Al-Azraq and surrounding basin sediments. Chapter Four presents the sampling of sediment sections for further micro-facies investigation based on initial core description and smear slides. The initial visual observation identified three sediment sections with observable laminations for further analysis. Discussed are the multiple methods applied to provide a robust interpretation of the origin and nature of these sediments. These analytical methods included: mineral grain size, X-Ray Florescence (XRF), Scanning Electron Microscopy (SEM) and Energy Dispersive Spectroscopy (EDS), and Magnetic Susceptibility (MS). Chapter Five reports the results of the analytical methods for each of the three sediment sections. Chapter Six discusses each of the sediment sections and the evidence for laminated sediments. Also discussed are the implications for paleoenvironmental variability. Chapter Seven presents a consideration of the evidence discussed supporting laminated sediments and the conclusions of the study. This chapter includes implications for future environments within the basin and region.

2. PROJECT BACKGROUND

Introduction

Evidence for global change in paleoenvironment records commonly reflect coarse timescales, thousands to tens of thousands of years. Laminated sediments provide valuable paleoarchives of fine-resolution sediments as short as intra-annual, seasonal layers to daily tidal layers (Mastalerz et al., 1999). Lamina and their components indicate short-term, dynamic environmental conditions and processes. It is widely understood that laminated sediments occur under very specific geological and climatic conditions ranging from glacial to arid settings. Conditions for deposition are a function of hydrometeorological parameters and subsequent preservation include a mix of protected basin setting, meromictic water column, anoxic conditions, low surface disturbance, changes in paleoproductivity, etc. (van Geen et al., 2003; Ojala, et al., 2014). Laminated sediments occur in arid paleolakes such as the Dead Sea, Israel (Ben Dor, 2019); Lake Van, Turkey (Ojala et al., 2012; Stockhecke et al., 2012), and Qaran Lake in Faiyum Oasis, Egypt (Marks et al., 2017).

Laminated Sediments

Laminated sediments are very small scale, fine sediment layered horizons occurring in depositional settings and sedimentary rocks. They are very thin to ultra-thin depositional layers differing in color or composition, and ranging in size from centimeters to millimeters in thickness (Kemp 1996; Boggs, 2012). There are multiple types of laminated sediments.

Laminae

Laminae can have contrasting colors, composition, texture, structure and/or thickness resulting in distinct changes in depositional pattern. Minerals, macrofossils, or biogenic particles are all potential components that compose laminae. Laminae reflect changes in environment or climate conditions.

Rhythmites

Rhythmites are laminated sediments demonstrating obvious repeating patterns representing repeating and cyclic deposition. The repeating pattern is often alternating sediment types such as shale and limestone cyclothems in the Carboniferous strata of the Midwest (Archer, 1999; Joeckel and Fielding, 2018).

Organic rhythmites are commonly found in vegetated and temperate regions. Spring nutrients promote algal blooms such as diatoms. Silica cell walls encase diatoms and their remains preserve in the sediment after completing their life cycle. This makes diatoms an excellent proxy for past environmental conditions. As a result of the life cycle of the diatoms endogenic rhythmites form the lightly colored summer layer seen in some sections of the drives. (Zolitschka, 2007).

Varves

The term varve, first used by DeGeer (1912), describes a seasonal lamination of coarse (summer) and fine (winter) material deposited in glacial lakes. The original Swedish word varves just means a cyclic event or form regardless of periodicity (Middleton, 2003). The now widely accepted definition for a varve is as follows by Hambrey (2011) “*Annually laminated deposits or varves are rhythmites with annual periodicity: annual layers of sediment or sedimentary rock are laid down through seasonal variations that result from precipitation, temperature, which influences precipitation rates and debris loads in runoff.*” Varves are distinguished from rhythmic sediments because they are demonstrated to be annual or subannual deposits (Zolitschka et al., 2015).

Varves are described as having two or more seasonal laminae. A single laminae layer or pair does not necessarily reflect a true calendar year but more a pattern of regularly fluctuating and cyclic climate. To be truly called a varved, a laminae must demonstrate

annual or subannual layering. This can be determined by radiometric age determination or other source of seasonal indicator (Zolitschka et al., 2015).

Varves in lacustrine environments are often composed of repeating or rhythmic components. These layers often reflect long periods of stable, seasonal lake productivity and are a mixture including: plant remains, carbonates, clastic, diatoms and soft organic tissues (Anderson, Dean, 1988). These can produce a layer indicative of the season and climate at the time of their deposition (Stockhecke et al., 2012). These layers can also contain mineral grain sizes indicative of different processes. A coarse layer or layer of silts can represent spring or summer runoff of terrestrial minerals. Fine or clay layers represent fall or winter deposition of terrestrial elements.

Formation Processes

Laminae form in several ways and settings including glacial, tidal, and lacustrine environments. Formation processes are influenced by geological setting, climate and human impacts. These range in time and process scale from stable (geologic) to highly variable (human impacts) (Figure 2.1.). Allochthonous minerals and organic detritus are washed into the basin from the surrounding watershed, or deposited through atmospheric circulation. Autochthonous generation of materials from biologic productivity, oxidation, or mineral precipitation create sediments in situ to the basin. Once within the basin multiple processes can alter the components further both before and after deposition through mixing, bioturbation, and diagenesis, etc.

Analyzing laminated sediments reveals the hydroclimate conditions at the time of deposition. Information about past environments comes from allochthonous materials from the watershed and atmospheric deposition. Autochthonous materials inform about lake processes within the basin. All sources are mixed within the water column and deposited as a lamina. (Hardy et al., 1996; Ojala and Alenius, 2005; Cockburn and Lamoureux, 2008).

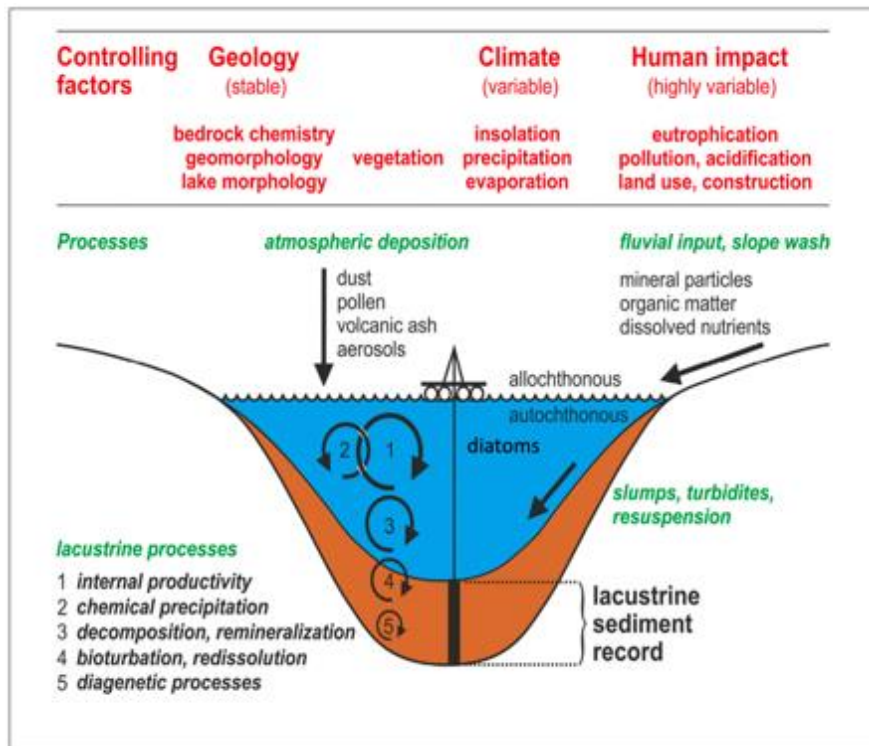


Figure 2.1. Processes of influencing formation of lacustrine sediment records including laminated sediments. These include controlling factors such as bedrock geology, structure, and tectonics; climate and human impacts ranging in scale from stable to highly variable. Modified from Zolitschka and Enters, 2009.

The laminae have contrasting colors, composition, texture, structure and/or thickness. The rhythmic pattern reflects the seasonal hydrological cycle in the environment of laminae deposition. The rhythmic nature of the deposits also represent the events in the environment occurring at the time of deposition such as changes in rainfall, temperature, algal bloom or even a dramatic event in the region such as an earthquake or flood.

Saline lakes form in arid regions where rainfall exceeds evaporation. Laminae of evaporite minerals precipitate as sediments during warmer phases. Evaporite minerals include halite, gypsum, anhydrite, calcite, dolomite, and other carbonates (Reeves, 2008). A high content of dissolved calcium carbonate can lead to the formation of a carbonate layer in the warmest months. The precipitation of the carbonates can be inorganically or biogenically

induced. With the rise of water temperatures or biogenic factors such as photosynthesis, dissolved CO₂ decreases, CO₃ increases and becomes saturated. Laminae can also form during cooler months, when there is little or no fluvial born sediment coming into the basin, clay-size particles settle to the bottom of the lake/basin, along with the remains of the planktonic organisms that flourished during the summer months. As a result, winter sediment is rich in clays. The clays will appear darker due to a small amount of carbon deposited in some. Together these laminae represent and seasonal and potentially subannual deposit reflecting warm and cool periods.

Varve formation

Varves are formed by seasonal variations and recorded in the sedimentary archives variations in hydroclimatic cycle. There are several models of formation processes which can be climate or basin specific. Organic varves form under anoxic conditions which preserve a range of organic source materials such as: pollen, plant debris, algae, humic, and fulvic acids (Sturm, Lotter, 1995; Kirkland, 2003). In one organic model spring and summer laminae are composed of several species of centric diatoms and calcite crystals. The fall season is composed of diatoms, Chrysophyte casts, fine clays and calcite. In this model spring/summer is light in color and winter is dark (Fig. 2.2.).

Clastic varves are formed by seasonal inputs of terrestrial mineral matter during spring flooding. The larger particles are washed in by meltwater and flash flooding events which also contain organic material (Sturm, Lotter, 1995). There is an upward grading of coarse to fine sediment layers which represent the seasonal changes from the more active spring summer to lower sedimentation in summer. The winter months are very low energy and sedimentation is in situ clays and fine silts (Fig. 2.2.)

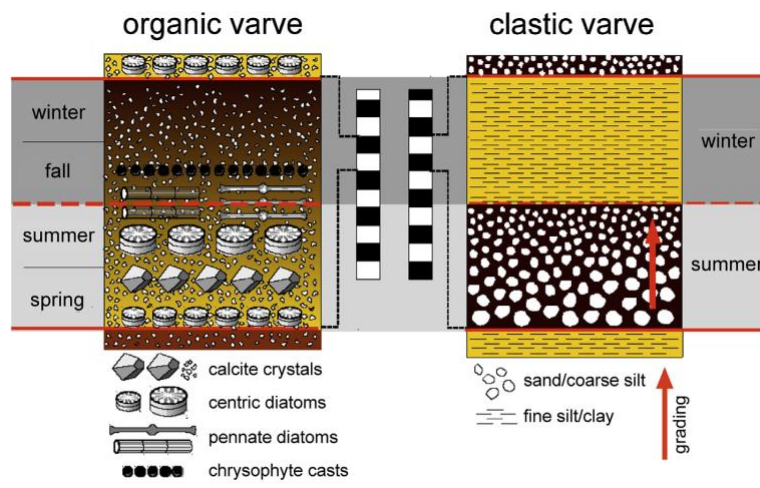


Figure 2.2. Simplified models for depositions of a calcareous and organic (mixed) varve and a clastic variation varve (modified from Sturm and Lotter, 1995).

Laminated Sediments in Arid Paleolakes

Laminated sediments have long been observed in arid to semi-arid environments across the globe. The model of seasonal deposition of these sediments originates in Scandinavian glacial deposits (deGeer, 1912), but an arid lake model of seasonal deposition comes from the famous finely laminated sediments along the Dead Sea's uplifted coastal margins.

Dead Sea Varve Debate

Dead Sea varves first described in the literature by Bloch (1944), developed the concept of seasonal white and dark couplets representing one annual cycle. These sediments became the classic example of varve deposition in arid environments for researchers for decades. A new debate surrounding the annual nature of the Dead Sea varves arises from three recent studies. Following the International Continental Drilling Program's new long core from the Dead Sea, the results published by Ben Dor et al. (2019) state the Dead Sea is the only deep hypersaline lake in the region known to have deposited long sequences of finely laminated, annually deposited sediments of varied compositions, including aragonite, gypsum, halite, and clastic sediments. They develop a new and more refined model of varve

formation identifying three types of laminated sediments: alternating aragonite and detritus (*aad*) laminated detritus (*ld*), and layered halite (*lh*). The *ld* laminated sediments containing pollen, represent flood deposits and are not annual. The *lh*, layered halite, are determined to be annual deposition representing summer precipitation of halite and winter flood materials. The *aad* laminae are also interpreted as annual deposits with the aragonite as summer evaporite deposition and the detrital laminae from winter rainy season (Haliva-Cohen et al., 2012).

The first argument against annual deposition, a very detailed pollen and grain size study by Lopez et al. (2016), demonstrates pollen was transported into the basin from the watershed and from reworked soils. Identified as flashflood events, these are not annual deposition. Detrital and aragonite laminae were shown to both be deposited during the rainy season, and therefore not representative of separate seasonal layers. Additionally aragonite is shown to precipitate in wet as well as dry seasons and therefore not exclusively an indicator of dry season evaporation. They conclude the laminated sediments from the basin marginal deposits of Ze'elim for the Holocene are not annual varves, and call for older laminated sediments to be reevaluated.

In a response to Ben Dor (2019), Bookman (2020) further argues the Dead Sea varves originally based on an observed “whitening” event; these events are rare and therefore not annual. Micro-facies analyses indicate detrital material to be from both single flood episodes and accumulation from an entire rainy season, also indicating these laminae are not annual. Ben Dor et al. (2018) concluded when these detrital episodes were capped with aragonite, whether single or multiple it represents an annual cycle. Bookman cites Lopez et al. (2016) as demonstrating aragonite deposition can be in both dry and wet season. The limitation to aragonite formation is not temperature dependent, but limited by availability of TCO₂. This

also argued against annual sedimentation and calls for reassessment of the varve characterization of Dead Sea laminates.

Qarun Lake, Egypt

An additional recent study of an arid lake with laminated sediments in the Eastern Mediterranean is Qarun Lake. It is located in the Faiyum Oasis, west of the Nile River in Egypt and contains a Holocene record of seasonally laminated deposits (Marks et al., 2017). While not nearly as deep or large as the Dead Sea, Qarun lake contains a fine resolution record of hydrologically driven sedimentation. It is also under the influence of the Eastern Mediterranean climate regime fluctuations. The Faiyum depression is one of the most important hydrological basins of the western Sahara, and similar to the Dead Sea is an endorheic, closed basin. It has throughout the Holocene been influenced periodically by inflows of flood waters from the Nile River. In addition to geochemistry, the sediment record is notable for its diatoms assemblage.

3. STUDY AREA

Introduction

The Jordan Plateau is between 35°30'-37°30'E latitude and 29°30'- 32°00'N longitude with the Al-Azraq basin covering an area approximately 708.9 km² (Ibrahim, 1996). The Jordan Plateau defines the two major drainages areas in the region, one channels water west into the Dead Sea Valley, a small extension of the East African Rift Valley; and the other flows east into the Al-Azraq basin in the desert interior (Burdon, 1959; and Bender, 1974). Jordan Plateau elevation ranges from 1734 to 500 meters at the Al-Azraq central Qa or mudflat area.

Oriented along the NW-SE axis of the Wadi Sirhan fault, the Al-Azraq basin is an endorheic or hydrologically-closed basin. Today it is a seasonally dry basin. Although the weather on the Jordan Plateau today is arid to semi-arid due to the rain shadow created by the western escarpment, the past environment was moister as demonstrated by the presence of lacustrine deposition within the basin (Davies, 2005; Ahmad and Davies, 2017). The Al-Azraq basin presently receives source water from winter rain and can remain without rain for several years. The Al-Azraq cored sediments are the focus of this research into the presence and characteristics of laminated lacustrine sediments.

The Al-Azraq basin has long been a critical source for water in the region (Ibrahim, 1996). Due to spring waters it has also long been considered an important trade and migration route throughout human history. Archaeological evidence dating back to the paleolithic demonstrates the basin's importance as a regional water source and migration route in the Pleistocene (Khoury, 2003; Davies, 2005). Previous research indicates a lacustrine environment associated with a moderate climate, and periods of high moisture (Davies, 2005; Ahmad and Davies, 2017). The capital city of Amman, the surrounding cities, and refugee

camps are within the Al-Azraq basin drainage and it remains a critical source of water today (Deutsche Welle, 2016; Hegazi, 2016).



Figure 3.1. Google image of the AL-Azraq Basin, Jordan. Inset shows the location of the AZ2 sediment core drill site on the eastern margin of the basin.

Geologic Setting

The principal geology of the region consists of Cretaceous and Tertiary chalky limestone. To the north and northeast of the basin are Oligocene basalts (Bender, 1974). The Dead Sea Transform is a dominant geotectonic feature in the area. The plate boundary covers an area between the Arabian plate and the Sinai microplate, and is a part of the African plate. These plates are part of the larger Syrian African Rift system covering 6000 km, which runs from central Africa, through the Red Sea, to the Taurus mountains in Turkey (Girdler, 1990; Klinger et al., 2000). There are several faults in the area that form the Jordan plateau and later the Azraq depression. The Jordan plateau consists of several fracture zones, the Karak-Feiha Fault which runs to the southeast. The Sirhan Fault extending from the Wadi Sirhan in Saudi Arabia in a NW-SE direction. The Fuluq Fault, the Karak-Feiha Fault and the

Sirhan Fault form the Hamza Graben forming the Al-Azraq basin (Margane et al., 2002; Kaudse, 2014).

The Qa Al-Azraq or central mudflat is bounded by the northern (AL-Bayda Fault) and southern (Ar-Rattam Fault) transform faults intersected by western (Al-Qaislyeh Fault) and eastern (Baqawiyya Fault) parallel faults. This rhomboid area forms a grabben whose downthrust block forms the Azraq depression, which is the deepest area of sediment deposition.

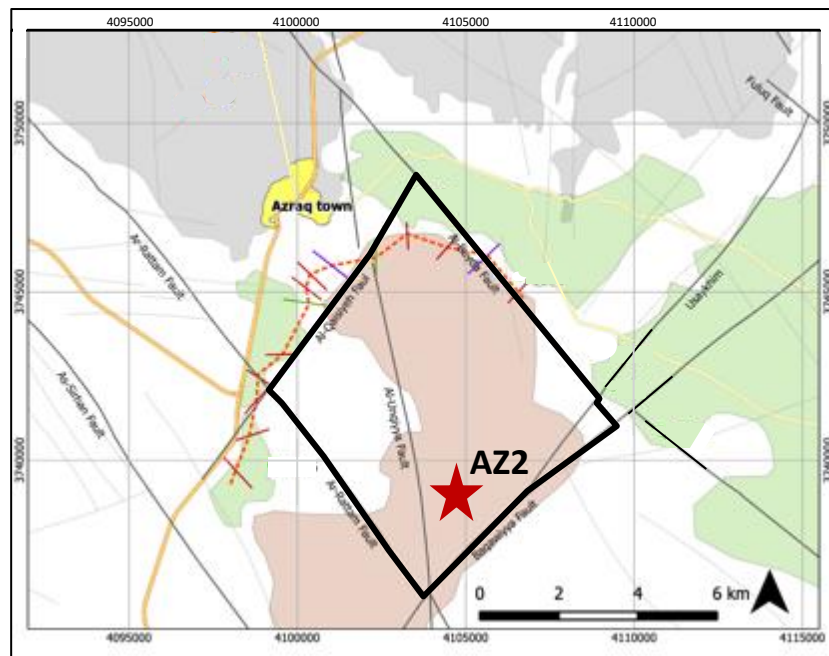


Figure 3.2. Al-Azraq basin illustrating the northern basalt margin (gray), the fault lines forming the basin; the Ar-Rattam Fault (southern) and AL-Bayda Fault (northern) transform faults intersected by Al-Qaislyeh Fault (western) and Baqawiyya Fault (eastern) faults. The coring location of AZ2 is indicated by the red star (modified from Kaudse, 2014).

Al-Azraq Basin and Sediments

The Al-Azraq basin is a closed basin system. These systems are found throughout arid and semi-arid regions of the world, where evaporation can exceeds precipitation (P/E) for periods of the year. Several wadies, tributaries, and a low depression covering a total ca.12,000 km² comprise the surface inputs to the Al-Azraq basin. The basin has several

springs fed by groundwater creating the Azraq Oasis. (Nelson, 1973; Garrard, 1998). There are two wetland areas, the wetlands to the north cover approximately 1.8 km², and the marshes to the south covering 5.6 km² (Nelson, 1973). The basin sediments are known to contain diatomite identified by Ala'li (1993).

The tectonically formed basin allows for a much deeper depression than is typical for playas in the eastern Al-Nafud Desert. The Al-Azraq depression perseveres a long and unique sedimentary archive of paleoenvironments. The sediments are primarily Quaternary age limestone sands, silts, clays, and evaporites.

Climate

Three climate characteristics define the Mediterranean climate regime characterizing the Jordan Plateau. First, the Subtropical High-Pressure belt is responsible for summer high temperatures along with a lack of summer precipitation (Lamb, 1995). Amman has a temperature ranging from 8.5°C in winter to a relatively high 26.5°C in summer. Second, the winter rainfall influenced by fluctuations in the Northern Jet Stream brings moist, cool air carried eastward across the Mediterranean Sea by temperate latitude westerlies brings the winter rains (Lamb, 1995). And third, the western escarpment creates a rain shadow resulting in the semi-arid to arid climate of the eastern desert. As a result, the eastern portion of the Jordanian desert receives less than 50-100 mm of rainfall annually (Rahman et al., 2015) (Fig 3.3). Current precipitation in the Al-Azraq region is low, less than 100 mm annually, but the Jebel Druze, Syria at the northwestern edge of the basin has been known to have as much as 200 mm/yr. (Nelson, 1973; Garrard and Byrd, 1992).

The Al-Azraq interior drainage flows into the basin which is an endorheic or hydrologically closed system. The closed basin provides a critical aquifer for the region. The aquifer has filled over geologic time spans during periods of higher than present precipitation. The basin's main source of water is from winter rainfall and may not receive any measurable

rain for several years (Al Naber, 2016). Presently, the people of the capital Amman, city of Zarqa, and several refugee camps depend on the aquifer and it is the main water supply for a majority of the country (Frances, 2015; Deutsche Welle 2016; Hegazi, 2016).

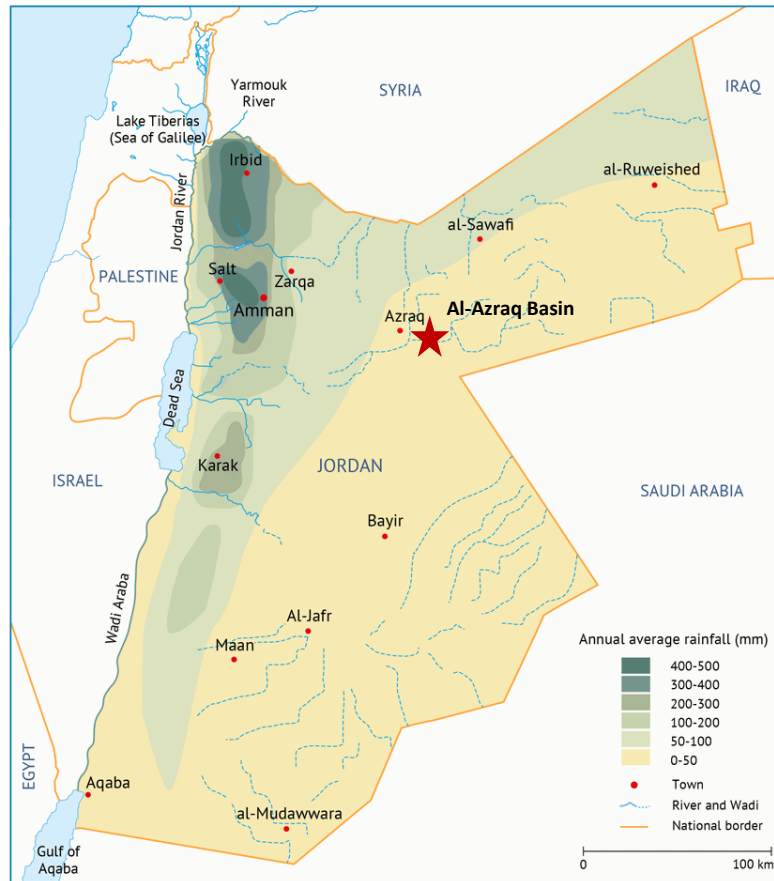


Figure 3.3 Isohyet map of Jordan with annual average rainfall. Modified from Jordanian Metrological Department Data, 2014.

Al-Azraq Basin Chronology

Dating of the Al-Azraq basin cored sediments comes from radiocarbon and infrared stimulated luminescence (IRSL). While radiocarbon and IRSL ages from samples in the upper sediments are not in agreement, they all return Pleistocene ages (Cane, 1972; Davies 2005). The IRSL ages however are internally consistent and reveal very old ages of the upper 11 meters of sediment. Sediments at greater depth are beyond the limits of IRSL.

Additional ages from sediment sections along the eastern basin margins include

radiocarbon ages, from 9,610 ka to 23,980 ka, OSL ages from 2.7 ka to 62 ka, and Uranium-Thorium (U-Th) series ages ranging between 93 ka and 118 from ‘Ayn Qasiyya (Jones and Richter, 2011; Cordova et al., 2008; Cordova et al., 2013). U-series age from a deposit at the outer margins near Umari 43 km to the south returned a much older age of 330 ka (Abed et al., 2008). Turner and Makhlouf (2005) also provide an age for exposed lacustrine sediments from Umari dating approximately 600 ka.

Table 3.1 Chronology of the Al-Azraq basin sediments.

Depth	Age	Method	Reference
5 m	11,460 ± 40 B.P.	charcoal	Davies, 2007
6.30 m	40,000 B.P.	charcoal	Cane, 1972
0.20-0.28 m	24.2 ± 2.0 ka	IRSL	Davies, 2005
5.37-5.46 m	163.3 ± 11.7 ka	IRSL	Davies, 2005
11.62 m	> 250 ka	IRSL	Davies, 2005

4. METHODOLOGY

The following describe the methods and procedures followed to prepare samples for multiple proxy analyses used in this study. Sample preparation followed protocols from the Limnological Research Center, Core Facility at the University of Minnesota (Schnurrenberger et al., 2003).

Initial Core Description

Initial core description (ICD) are the observed lithologic characteristics of the cored sediments. After the core is split and the interior surface cleaned, sediment characteristics are described. These include: grain texture, fabric, bedding structure, color (based on the Munsell color chart system), inclusions, and any features or marks. Images are taken of sections following surface cleaning and lithological description. The ICD also notes any known incidents occurring to the core between the points recovery and examination. This is beginning process for all sediment analysis.

Smear Slides

After the routine examination and description of cored sediment, smear slide sampling at regular intervals provides a quick micro view of sediment composition. A smear slide is a small drop on the end of a toothpick of unprocessed sediment mixed with water and covered for visual observation at 10x. This is a quick and efficient method for planning more in-depth, high resolution sampling and processing, which is both time consuming and expensive. Smear slides are a quick method for examining detailed stratigraphic changes, changes in grain size, changes in color, basic mineralogy, and any organic inclusions within the sediments. Fifty samples processed for smear slide examined sediments for potential laminated horizons.

Grain Size

Grain size analysis in paleoenvironmental research indicates relationships between transport energy and grain particle size at the time of deposition (McCave and Syvitski, 1991). Processing sediments for grain size analysis also reveals how much or how little organic matter is present in the sediment. Grain size analysis provides the percentages of clay, silt, and sand by particle diameter in the sediment identifying the characteristics, frequency, and distribution of sample sediments. Grain size also contributes to determining the mineral source of sediments, as well as particle size range- clay, silt, and sand. One hundred and eighty-four sediment samples run on the Coulter Laser Particle Analyzer LS2000 in the Department of Earth and Environmental Sciences at the University of Missouri Kansas City identified particle sizes after dissolution of organics and carbonates. Nine replicates were measured for each sample.

Statistics on grain size data provide a range of metrics for determining energy environment parameters. The median identifies sediment midpoint of the distribution between coarser and finer particles. Mean refers to the arithmetic average grain size. Calculating the standard deviation determines the sorting values of sediments. Kurtosis is the degree of peakedness or flatness related to the probability distribution relative to benchmark normal distribution. Skewness measures the distribution within the tails of a curve and is independent of the sorting of the samples (Folk and Ward, 1957).

The Shepard/Folk are classification systems used to characterize sediment distributions. Folk's (1974) classification system uses the term mud instead of silt and is based on two diagrams. The Folk system has a less detailed separation between clay and silt. The major difference between the Folk (1974) and Shepard (1954) system is the Folk system focuses on the velocity of a current at the time of deposition and the maximum grain size of debris present (Poppe et al., 2013). Shepard's system is more focused on the ratios of clay,

silt, and sand making it a more detailed system in regard to sediment classification. (Poppe et al., 2013). The USGS program SEDCLASS helps to standardize the classification systems and is used in this research (Poppe et al. 2003).

X-Ray Florescence

X-Ray Fluorescence spectroscopy (XRF) is a non-destructive analysis providing qualitative and quantitative data through the emission of secondary X-rays. When a primary X-ray source bombards a sample with high energy X-rays or gamma rays, samples emit secondary X-rays. The emissions or fluorescence characterize specific elements and their concentrations are based on count rates or energy intensities (Fitton, 1997). This provides elemental and geochemical data. (Richter et al. 2006).

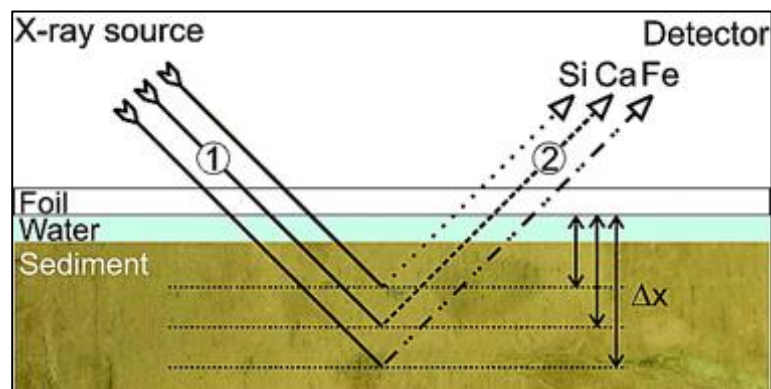


Figure 4.1. X-Ray Fluorescence Analysis schematic. The x-ray beam ionizes elements (1); radiation emits element-specific fluorescence (2), registering on the detector. Heavier elements emit stronger fluorescence energy (Richter et al., 2006).

Scanning core sediments through the XRF beam produces an element map. XRF run at The Scripps Institution of Oceanography analyzed sediment sections every 0.5mm and 1mm, at 10 seconds at 10kv, 20 seconds at 30 kv, and 30 seconds at 50kv, on a total of 176 cm. Cores lightly scraped to expose less oxidized surface layers of sediment were then covered with 4 mm thick SPEX Certi Prep Ultralene VR foil to prevent excess drying or contamination of the sediments during analysis as the XRF sensor moved downcore. Element mapping of laminated sediment generates positive and negative spikes in elements

corresponding to visible changes in light and dark sediments or changes in element concentrations.

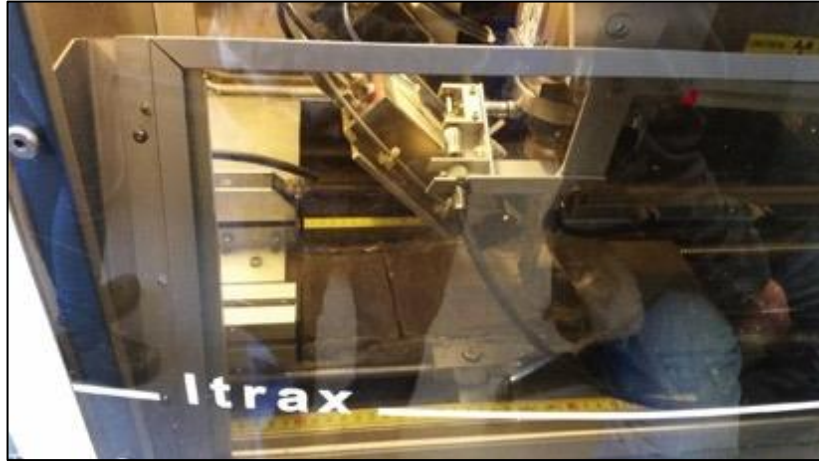


Figure 4.2. Photograph of XRF Itrax scanning sediment Section 3.

Scanning Electron Microscopy and Energy-Dispersive X-Ray Spectroscopy

The scanning electron microscopy (SEM) produces a signal from a high energy beam of electrons focused on a sample which produces an image from the position of the beam and element specific intensity of the detection pattern. When atoms and electrons interact, the signals produced relay information about the sample's surface topography. The Energy-Dispersive X-Ray Spectroscopy (EDS) provides a distribution of elemental composition (McMullan, 2006). SEM together with EDS analysis provide nano-scale information of specific samples and the elemental composition.

Preparation of samples for SEM analysis includes either coating with a thin layer of conducting material, commonly carbon, gold, or some other metal or alloy, or simply run without any coating in low vacuum mode (Goldstein et al., 2003). Coating materials, referred to as sputter coating, such as metals can increase signal-to-noise ratio during imaging due to their high conductivity, thereby producing better quality images. Coating can eliminate electron traps at the surface of non-conductive materials. Twelve AZ2 samples were scanned on Tescan Vega 3 LMU variable pressure SEM equipped with a Bruker Quantax Energy-

Dispersive Spectroscopy (EDS) system and a Tescan color cathodoluminescence detector. A majority of samples had no coatings, a few samples had carbon coatings. As part of the SEM analysis, Energy-dispersive X-Ray Spectroscopy (EDS) were performed on the samples to assist with the analysis of elements in the sample.

Table 4.1. Types of analyses performed on AZ2 cored sediments, instrumentation, laboratory, and number of samples processed.

Analysis	Instrument	Laboratory	# Samples
Core description	-	UMKC	25 meters
Smear slides	-	UMKC	12
Grain Size	Coulter Laser analyzer LS200	UMKC	184
XRF		Woods Hole Oceanographic Institute & Scripps Institute	176 cm (scanned)
SEM/EDS	Tescan Vega 3 SEM/Bruker Quantax EDS	UMKC	12/5
Total Samples			414

5. RESULTS

Sample Selection

AZ2 cored sediments extend from 6.5 m to 36 m below the surface. There is a gap between 25 and 35 m where during drilling a 10 m thick bed of white quartz sand cause a blow out of the hole and no sediments were recovered. Initial core description, smear slide and grain size analyses, identified sections of laminated sediments for further analyses.

Initial Core Description

The initial core description (ICD) of the total 36 m of AZ2 sediment core and photographic images revealed a detailed lithology of intricate fine sediment deposition with many subtle changes in texture and coloring. On the basis of this visual inspection, three sections totaling 4.5 m of sediment from the original 36 m exhibited laminated horizons of both fine sediments and alternating colors (Figure 5.1). These sections, referred to here as Section 1 (12.5 m to 13.80 m), Section 2 (14 m to 15.5 m), and Section 3 (20.5 m to 22 m), are the focus of the remainder of this study.

AZ2 sediment lithology 35 is a sandy, silty mix with visible lens of clays, carbonates, and oxidation throughout the core. The core is well laminated in sections. Between 11 m and 20 m there are repeating lens of oxidation along with thin dark and light laminated sediments. There is visible slumping, rounded and embedded clumps of sediment from upper sections. There are also numerous angular slumping of horizons and some vertical faulting. These are all possibly due to tectonic events starting at 14 m. The core has some missing and broken sections, but it is largely intact. The sediments have white spots and horizons are possible carbonate and halite deposits. There is a 10-meter section between 25 m and 35 m of quartz sand bed lost during drilling. Of the 25 m recovered and examined, ten sections were pulled for further analysis and of these, three underwent micro-facies analysis of laminations. These three sections are the focus of this study.



Figure 5.1. AZ2 cored sediments images, lithology, and legend. (a.) 11 m to 22 m exhibited laminated sediments; (b.) Section 1 (14 m to 14.50 m, Section 2 (19.00 m to 19.50 m), Section 3 (20.5 m to 22 m) analyzed for micro-facies.

Section 1

Section 1 extends from 14.00 m to 14.50 m. The lithology of the first 1 cm was very dry friable, with faint white carbonate. Colors are stated according to the Munsell Color Chart designation (2009). 1.5 cm was dark olive-gray clay (7.5Y, 4/3-5/3) color, and 6-8 cm horizon of yellowish and olive gray (7.5Y, 5/2). 10-12 cm sediments were more brittle, hints of lighter gray 5/1 and rust colored horizontal streaks. 12-13 cm were olive yellow (7.5Y, 6/3), still brittle with possible halite. 13-19.74 cm were dark olive (7.5Y, 4/3) and brittle. 19-20 cm was horizon of oxidation. 19.75-21 cm was a yellowish-brown horizontal stripe (10YR, 5/6). 21-37 cm were grayish-yellow brown (10YR, 5/2) with hints of possible halite. 37-39.5cm were grayish-yellow brown (10YR, 6/2) smooth, hard clay and more possible halite. 39.5-41cm very dark grayish-olive stripe (7.5Y, 5/2). 41-44.5 cm were grayish olive with vertical white sediment (7.5Y, 6/2). 44.6-50 cm olive yellow (7.5Y, 6/3) with horizontal halite sediments. The section was 50 cm in length.

Section 2

Section 2 extends from 19.00 m to 19.76 m and overall composed of clay mixed with some silt. The top of the section, from 1 cm to 8.5 cm was friable. Colors were alternating thin layers of black (5Y 2.5/1) and a dark gray (5Y 4/1) horizons. There were also some yellowish red sediments (5YR 4/6) with yellow (10YR 8/6) streaks. There was a small crack running horizontally across the core at 8 cm to 9 cm. From 14.4 cm to 16.5 cm a chunk was missing from middle of core. At 11.5 cm there was a black line (5Y 2.5/1) running half way across the width of the sediments. 12 cm to 14.5 cm another thicker black line (5Y 2.5/1) ran across core and sloping down to 15.5 cm. There was a break in the core at 19 cm. At 22 cm extending to 22.5 cm was a crack sloping to the right. At 24 cm the color lightens to gray (5Y 4/1). At 32 cm small cracks with reddish/orange colors, suspected to be Fe, ranging from yellowish red (5YR 4/6) to dark grey (5YR 4/1). There was a small chunk missing at 33 cm

just under the possible Fe. Overall, color was mottled, dark gray (5Y 4/1) and black (5Y 2.5/1). More possible Fe at 38 cm. There was a break in the core at 39 cm on the bottom of the broken section there was dark gray (5YR 4/1) mixed with yellow (10YR 8/6). From 33.5 cm down to 42 cm color was yellowish red (5YR 4/6). The color from 43 cm to 48 cm was dark gray (5Y 4/1). At 49cm there was a missing chunk. At 53 cm a small crack. At 55 cm and 56 cm there were black (5Y 2.5/1) horizons, mottled with dark gray (5Y 4/1). At 61 cm to 65 cm the color returned to a solid black (5Y 2.5/1). At 67cm to 76 cm more black (2.5Y 5/1) horizons, with some mottled dark gray (5Y 4/1). The end of the core is uneven and somewhat friable. The sediment section measured 76 cm long.

Section 3

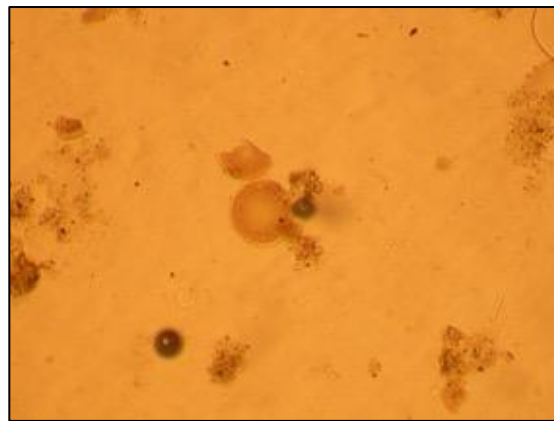
Section 3 extended from 20.50 m to 20.96 m. From 0 to 5 cm was friable with some yellowish red (5YR 4/6) showing in the broken chunks. The overall color of the sediments was a mottled black (5Y 2.5/1) and dark gray (5Y 4/1). The core is mottled from 0 cm to 15.5 cm, at 15.5 cm to 16 cm was streaking of yellowish red (5YR 4/6), then again at 17 cm streaks ran the width of the core. From 17 cm to 20 cm the core was a solid dark gray color (10YR 4/1). At 21 cm was a small break with a yellowish red streak (5YR 4/6). At 22 cm was a small line of yellowish red (5YR 4/6). At 27 cm a small black line (5Y 2.5/1) curved upwards. At 29 cm to 30.5 cm more yellowish red (5YR 4/6). From 40 cm to 46cm the sediment was a solid chunk mottled with black (5Y 2.5/1) and dark gray (5Y 4/1). After cleaning, photos, and being exposed to air it appears there is some oxidation at 21.5 cm and 28 cm to 30 cm. The upper section from 0 cm to 21.5 cm has turned a little darker than originally reported in field notes at dark gray (10YR 4/1). The sediment section is 46 cm long.

Smear Slides

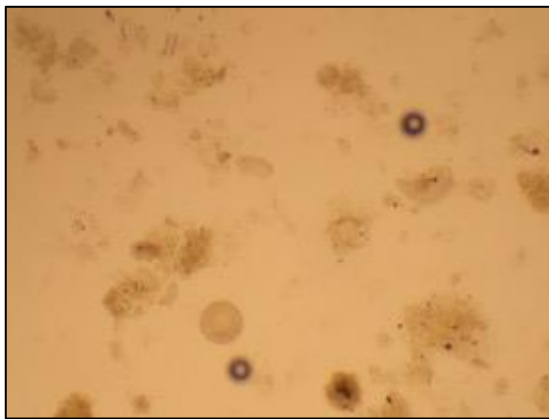
In addition to the visual observations in the initial core descriptions, smear slides identified diatoms, minerals, and other materials found in the sediments. It is also a relatively quick way to determine where further sampling should be focused. Twelve smear slides examined samples at the micro level for components of laminated sediments. This investigation revealed fragments and partial valves of two types of diatoms, one centric and one cylindrical. Also observed were algae and quartz sand grains.



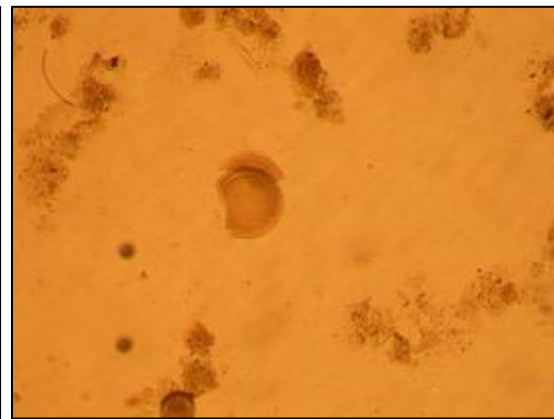
a. Section 1 *Stephanodiscus* sp.



b. Section 2 *Stephanodiscus* sp. and algae



c. Section 2 *Stephanodiscus* sp. and algae

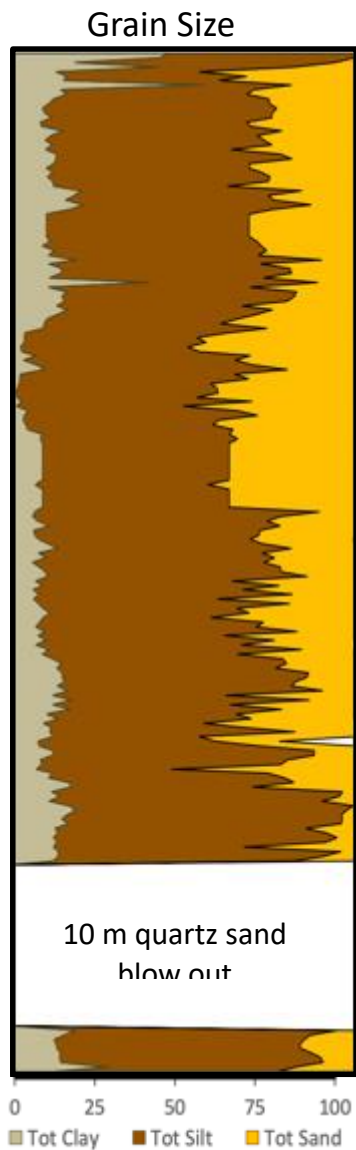


d. Section 3 *Stephanodiscus* sp.

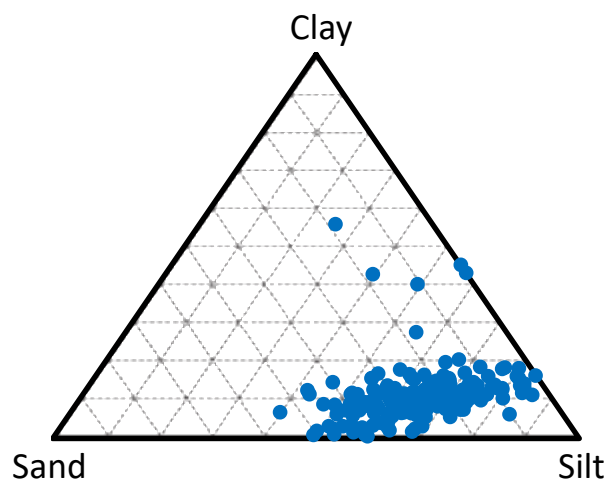
Figure 5.2. Smear slides images (a., b., c., and d.) from Sections 1, 2, and 3 showing the presence of the centric diatom *Stephanodiscus* sp., random clumps of algal material, and mineral detritus.

Grain Size

Grain size analysis is important for characterizing the energy environments of lacustrine sediments. The total percent clay ranges from a minimum of 11.20% to a maximum of 55.71% throughout the core, averaging 10.26%, with 26.23% in the uppermost 5 m. The total percent of clay is relatively constant, but total percentages of silt and sand fluctuate. Silt ranges from 25.85% to 85.33% and averaging 63% throughout the core. Between 25.24 m and 25.51 m silt averages 79% and peaks as high as 85.33%. Sand percent ranges from zero at the top of the core to 53.47%. Sand lenses occur at 11.49 m with 50% sand and 23.73 m with 53.47% sand. Lenses of silt, silty-sand, and clayey-silt punctuate the core. Above 25 m grain size distribution fines upwards and correspond to the formation and deepening of the lake. A 10-meter-thick quartz sand bed lies between 35 to 25 m. Below 35 m angular chert gravels were encountered, but not recovered (Figure 5.3).



a.



b.

Figure 5.3. Graph of cumulative percent grains size distribution of clay, silt and sand (a.); tertiary diagram of percent grain size distribution illustrating Shepard/Folk classification as predominately sand/silt throughout the cored sediments (b.).

Section 1

In Section 1 the grain size shows silt percent to be consistently averaging around 54.82% and sand percent averaging 36.74%. Total clay percent for Section 1 remains low at below 10%. Folk and Shepard classification of the sediment in Section 1 is Sandy Silt.

Section 2

In Section 2 the silt percent fluctuates more between 65 and 74%, with sand fluctuating between 15 and 38%. Total clay percent for Section 2 also remains low below 10% with the exception of one reading at 10.37%. Although there are increased fluctuations in silt and sand percentage, the Shepard/Folk designation for this section remains Sandy Silt. Tan colored and oxidized nodules occur throughout the section.

Section 3

In Section 3 there are far fewer grain size samples at only four. The silt percent ranges from 64 to 72%, and the range of sand percent between 13 and 22%. Clay percent is slightly higher ranging from 11 to 14.41%. This still earns a classification of Shepard/Folk Sandy Silt with the exception of 22 m which earns a Shepard designation of Clayey Silt.

X-Ray Florescence

The XRF analysis performed at Scripps Oceanographic Institute scanned 176 cm of sediment at 0.5 mm intervals and 10 kv. Scans included Section 1 from 14.30-14.46 m, Section 2 from 19.62-19.76 m, and Section 3 from 20.83-20.66 m show fluctuations in major (Al, Ca, Fe, Si) and trace elements (Cl, Cr, K, P, Mn, Ti, S, Sr, Rh) occur throughout the section sediments (Appendix C). Overall the micro-scale mm scanning reveals repeating patterns of fine resolution fluctuations.

Section 1

The XRF data, reported in Area and DArea for each major and trace element, were tested for accuracy of percent fit to model calculations by dividing the DArea model

prediction by the Area of the recorded element scan. Element scans are considered most accurate below 10% and not above 25% (Witter et al., 2018). The XRF pattern of major elements Aluminum (Al) and Calcium (Ca), particularly Ca, fluctuate in very regular, short nearly 5 mm periods, with very low amplitude over the length of the Section 1. Al demonstrates longer-term trends of possibly two oscillating cycles. There are two sharply decreasing excursions at 135 mm, 185 mm, and the very bottom in Fe, and Si patterns. Al has the decrease at 185 mm and gradual decrease at the bottom.. Between the downward excursions at 135 mm and 185 mm, both Fe and Si increase in threshold to new highs (Figure 5.4.).

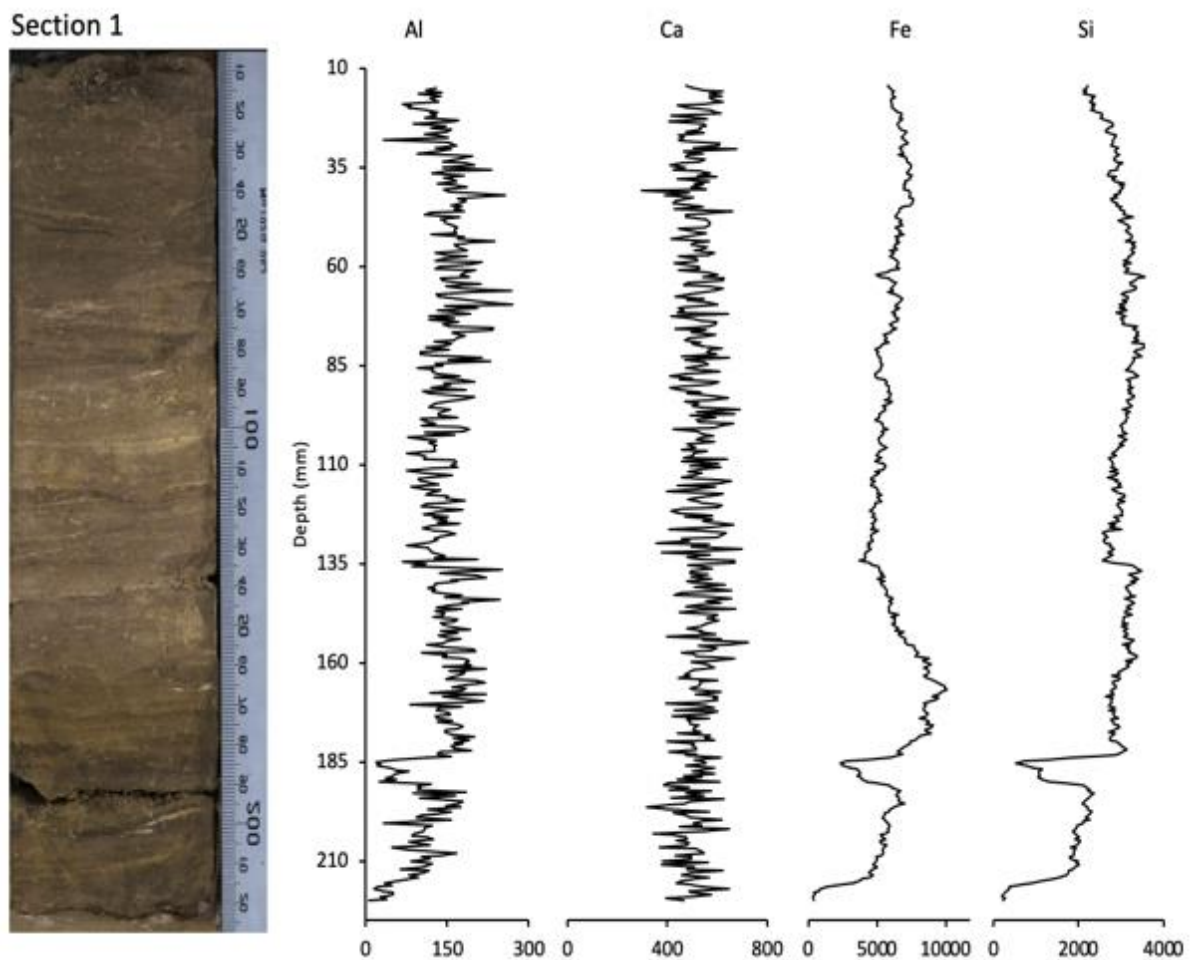


Figure 5.4. XRF graphs of major elements in Section 1. Scale in mm and XRF units measured in Area.

The Section 1 minor elements reflect similar fine mm scale fluctuating patterns seen in the major elements. Strontium (Sr) was recorded only Section 1. Chromium (Cr) very closely mirrors the low amplitude fluctuations of Ca and is in phase. for the entire length of Section 1. Overall patterns are harder to discern with the exception of Cl, K, Sr, and Ti. All decrease at 185 mm. Rh and Mn both increase slightly at 195 mm (Figure 5.5.).

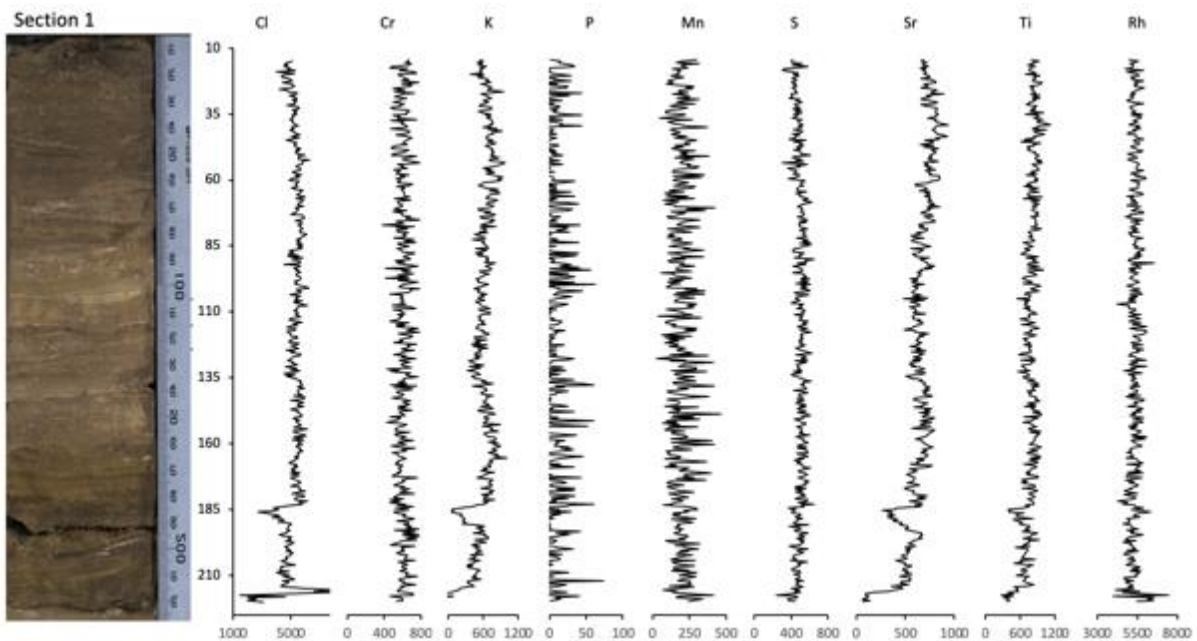


Figure 5.5. XRF graphs of minor elements in Section 1. Scale in mm and XRF units measured in Area.

Section 2

The XRF major elements in Section 2 have a very similar short period, low amplitude pattern as seen in Section 1 major elements for the upper 80 mm. From 80 mm to the bottom the pattern in Ca changes dramatically with two large episodes of increase over 20 mm which each increase the threshold to a new average. Al follows an opposite pattern of decreasing over the same range as the double peaks in Ca. Al is slightly out of phase beginning before 120 mm and extending to 60 mm. Fe and Si also show a similar trend with this long-term cycle in Al, but with less extreme fluctuations over the same range Al (Figure 5.6.).

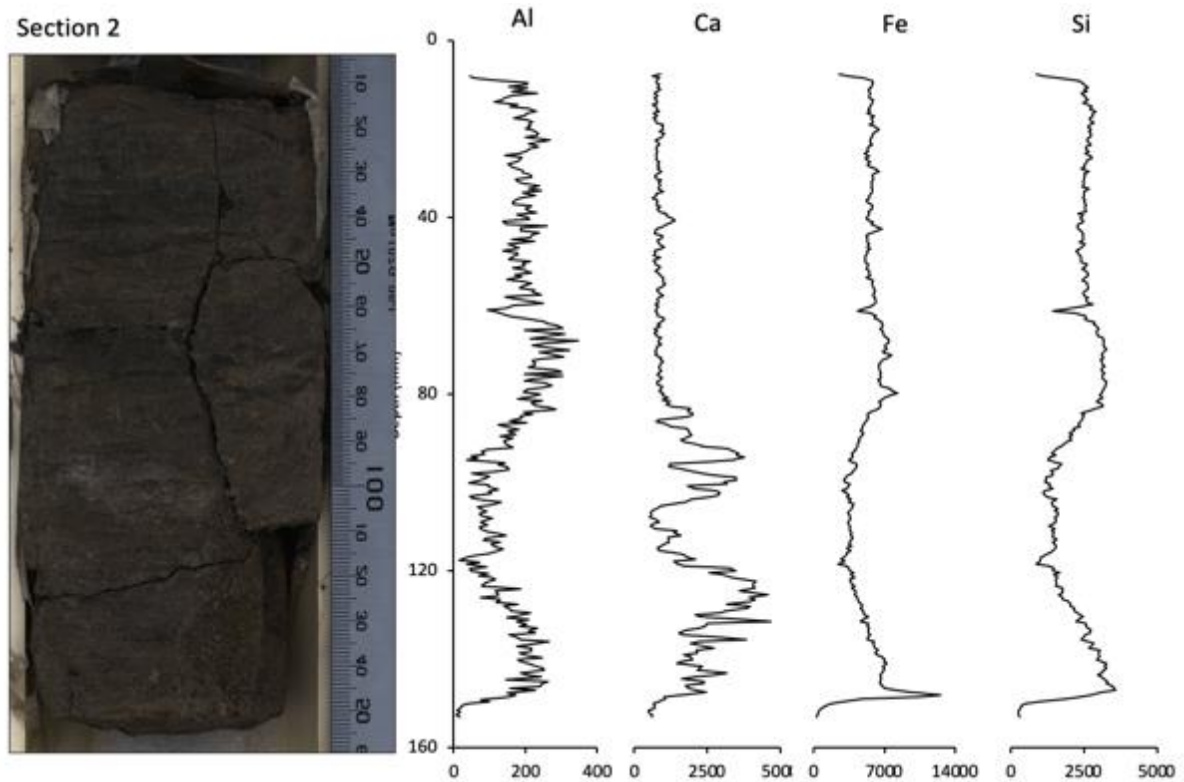


Figure 5.6. XRF graphs of major elements in Section 2. Scale in mm and XRF units measured in Area.

The Section 2 graphs of minor elements illustrate similar low amplitude fluctuations seen in Section 1 in Section 2 elements Cl, Cr, Mn, and P (Figure 5.7.). However, there is significant range in cycle trends in K, S, and Ti. Potassium (K) very closely mirrors the major element pattern of Al, and fluctuations in Ti are similar, but less intense than Al and K. All have cycles spanning approximately 150 mm to 60 mm. Sulfur (S) very closely mirrors the pattern of Ca with two large peaks below 60 mm.

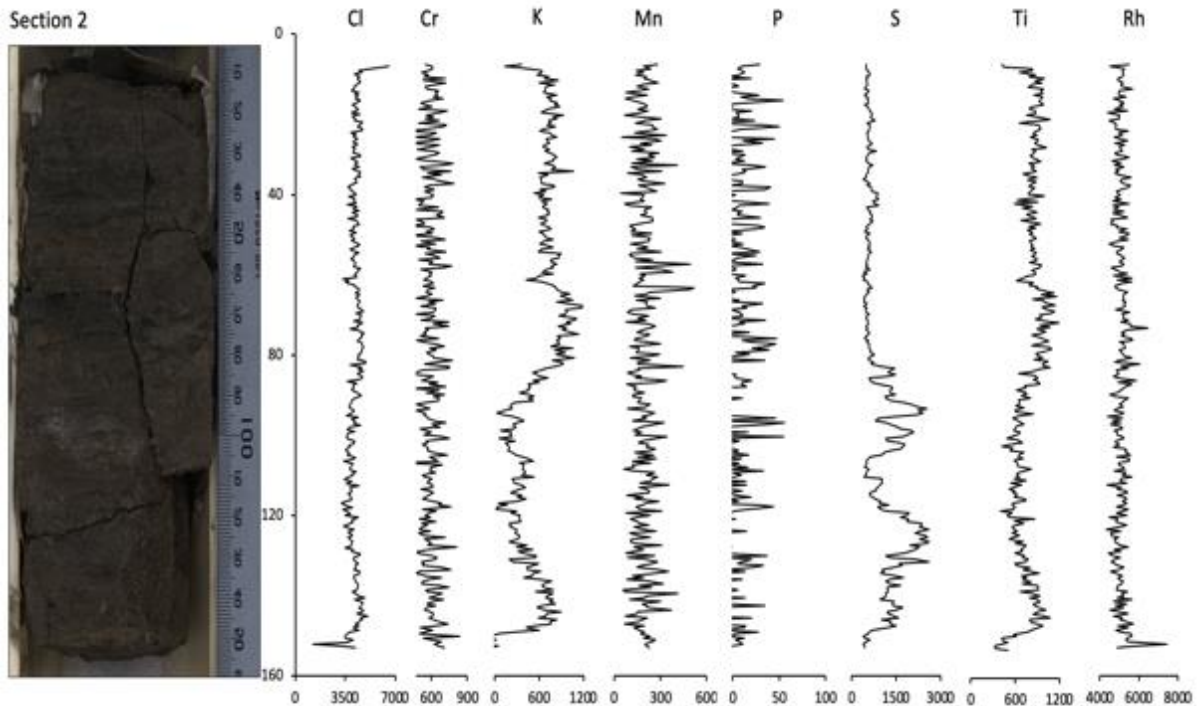


Figure 5.7. XRF graphs of minor elements in Section 2. Scale in mm and XRF units measured in Area.

Section 3

In Section 3 the major element of Ca for the majority of the section fluctuates very tightly around a narrow mean with four significant peaks occurring at: 3.5 mm, 122.5 mm, 148.5mm, and 158.5 mm. These peaks average 2886, with the highest peak extending to 5355 at 122.5 mm. For the remainder of the section Ca averages 74% less at 738. Al also fluctuate around a slightly less narrow mean, but the overall trend increases gradually in a bend from 160 mm to 60 mm. This same gradual increase and decrease is mirrored in Fe and Si, whose fluctuations are very low amplitude with very short periods. Fe and Si both have sharp decreases at 120 mm, concurrent with the very large spike in value in Ca.

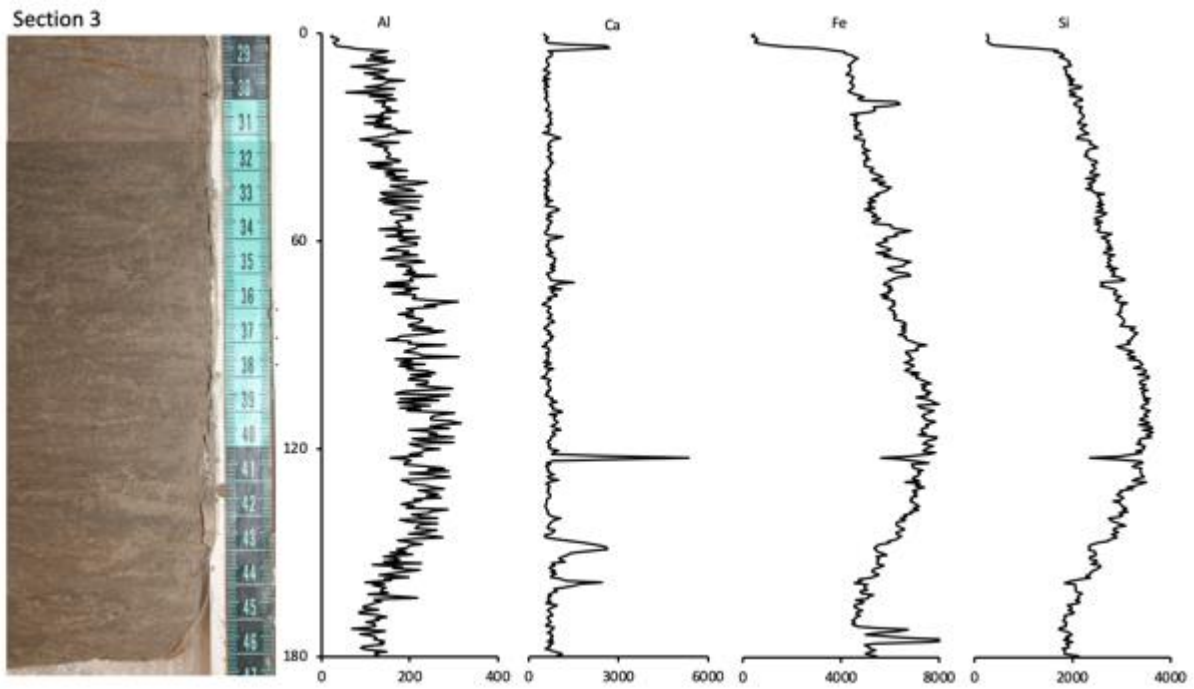


Figure 5.8. XRF graphs of major elements in Section 3. Scale in mm and XRF units measured in Area.

Some Section 3 minor element fluctuations mirror particular major elements. Chlorine (Cl) and sulfur (S) very closely mirror the very low amplitude fluctuations and very narrow mean of Ca, and each have four major peaks in values, including the large spike at 120 mm. Potassium (K) and titanium (Ti) also mirror the gradual increase and decrease of Aluminum (Al), iron (Fe), and silica (Si).

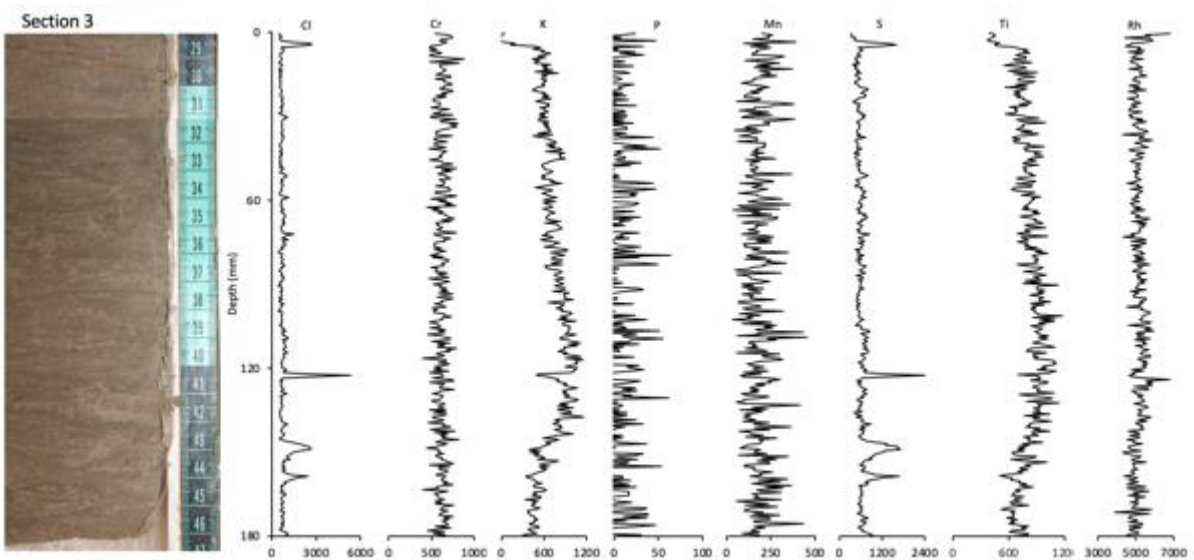


Figure 5.9. XRF graphs of minor elements in Section 3. Scale in mm and XRF units measured in Area.

Scanning Electron Microscopy (SEM)

Scanning Electron Microscopy (SEM) of specific representative samples from the three sections further explored the diatoms first observed in the smear slides and inferred from the higher amounts of silica (Si) detected by the XRF. The nano-scale focus and detail provided by SEM images allowed the identification of two species of centric diatom, one discoid and the other cylindrical. All sediment sections contain a mix of *Stephanodiscus* sp. and *Aulacoseira lirata* (Figure 5.10.). The images revealed diatoms species mixed together, stacked, crushed, and fragmented with a low number of intact frustules. Diatoms frequently observed in SEM images to be embedded or encrusted in a matrix.

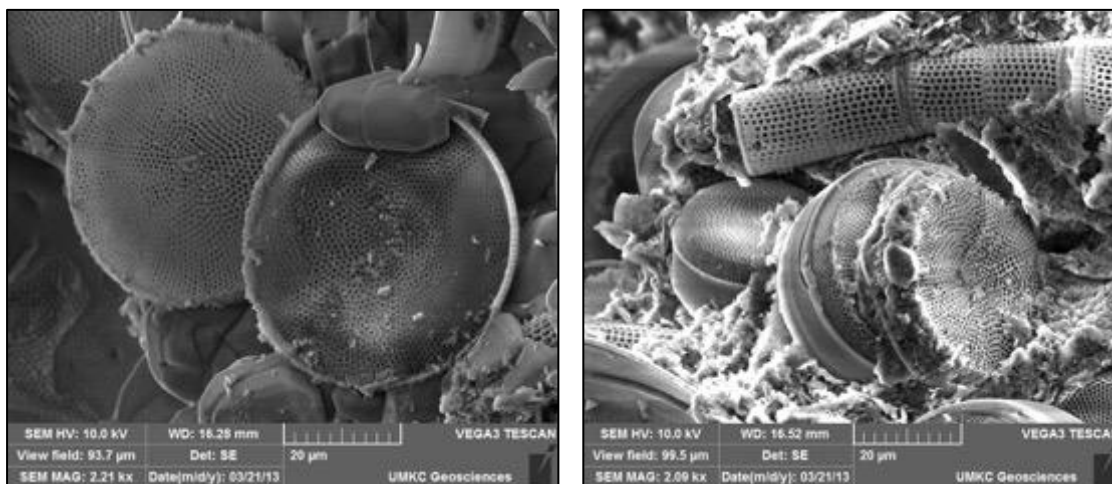


Figure 5.10. SEM images of diatoms from Section 3, identified as *Stephanodiscus* sp. (left) and Section 2, identified as *Aulacoseira lirata* (right).

Energy Dispersive Spectroscopy (EDS)

Energy Dispersive Spectroscopy (EDS) run on the same SEM samples found elevated levels of Al, as high as 30 cps/eV; Si and O at extremely high levels to 70 cps/eV. The EDS also recorded lower levels of Cl, Fe, K, Mg, and Na. EDS determined halite to be in all three sediment sections. Additionally, the EDS confirmed the composition of the diatoms with extremely high Si and O. analysis further identified the composition of the encrusting matrix

as halite (NaCl) and Fe-Mg-Al clays (figure 5.11.). A total of five samples underwent EDS analysis.

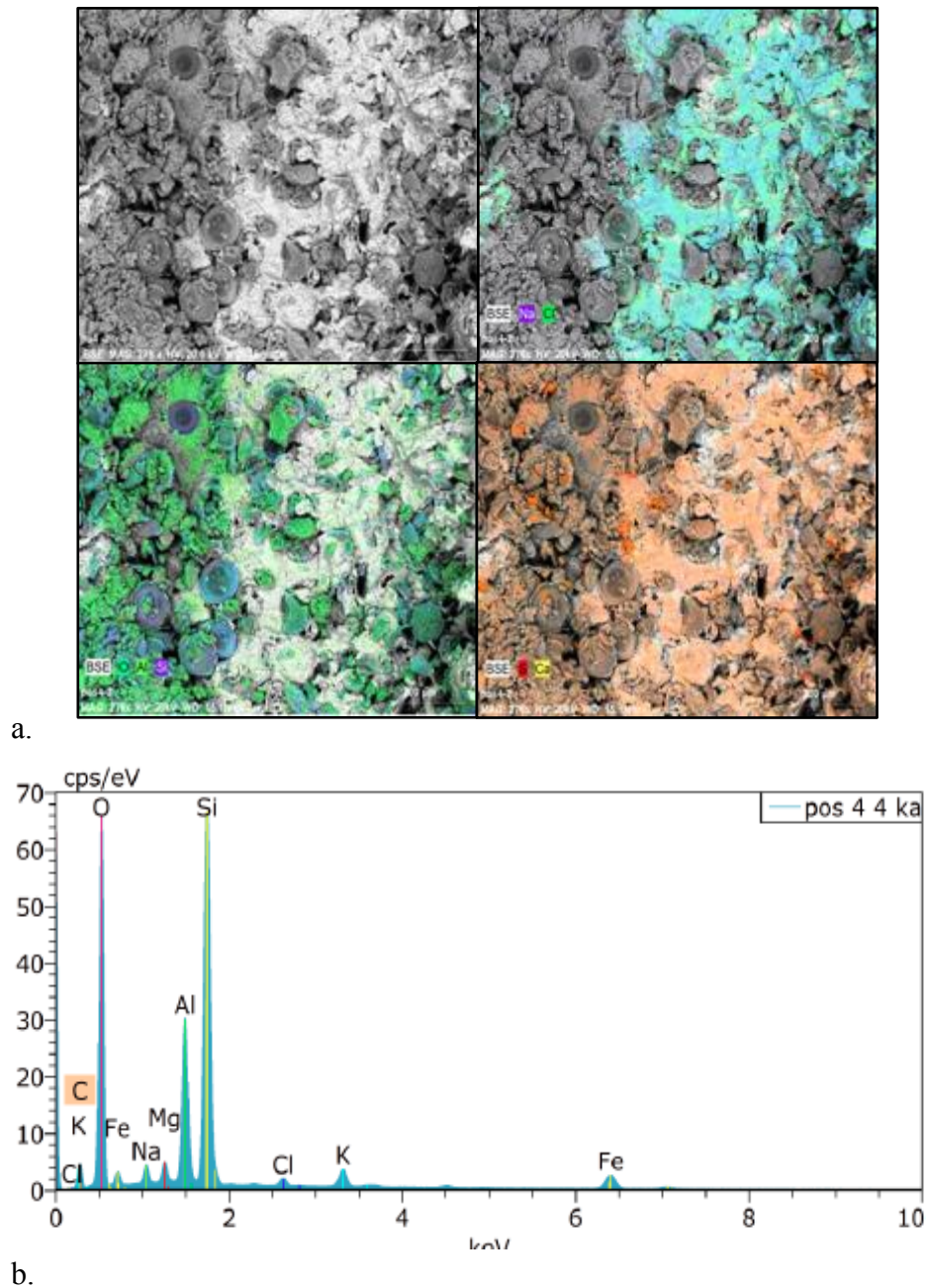


Figure 5.11. SEM image multiple centric diatom frustules embedded in matrix (a. upper left); EDS images (a. upper right, lower left and lower right) of diatom frustules highlighting dominant elements. Image identifying the encrusting matrix as halite (teal color) by the presence of sodium (Na) and chloride (Cl) (a. upper right). Image with green, blue, and purple identifies the diatoms with extremely high (Si, O) in a matrix of Al clays (a. lower left). Image with orange highlights identifies sulfur (S) and calcium (Ca) in matrix of crushed tests, Fe-Mg-Al clay (a. lower right). The EDS diagram further confirms element identification and abundances.

6. DISCUSSION

Introduction

This project investigates the cyclic nature of laminated sediments from the Al-Azraq Basin, Jordan. A 35 m sediment core was described, and analyzed with smear slides and grain size analysis. These initial analyses demonstrated the presence of very thin laminated sediments of alternating layers of light and dark sediment laminae with varying textures. The rhythmites consist of silt, clay, carbonate, clastic minerals, organic detritus, and diatoms. This is consistent with rhythmically laminated sediment (Glenn and Kelts 1991). Three sediment sections were chosen to further investigate the rhythmite layers through micro level geochemical analysis. X-Ray Florescence, performed at the Scripps Institution of Oceanography, scanned these sections at 5 mm intervals. Additionally, SEM and EDS analysis illustrated abundant diatoms which were identified to species.

Section 1

Calcium (Ca) is a proxy for carbonates which often form through evaporation during a dry season. Work published by Lopez et al. (2016), Belmaker (2019), and Bookman (2019) identify the formation of the evaporite carbonate aragonite as not linked exclusively to temperature, but is alkalinity limited. Until the specific source of the calcium in AZ2 sediments is identified through X-Ray Diffraction analysis, more cannot be determined on the seasonality of the Ca. However, the Ca relationship with other elements can provide some insights in to the nature of the laminated deposits. In all three sections Ca returns very fine scale alternating Ca

Silica (Si) is proxy for a mix of terrestrial silica and diatoms which are an *in situ* source of Si (Neuberger 2018). Both represent wet periods. The terrestrial clastic material washes in from the watershed during winter or spring flooding. The flood waters create an influx of fresh water bringing with it terrestrial silicas and organic detritus, and these in turn

supply the diatoms with silica and nutrients. In all three sediment sections Ca and Si alternate in very short and tight periods of low amplitude at the sub-millimeter scale which indicate short and alternating depositional conditions. There are also longer term trends visible in these patterns, as well as abrupt transitions in pattern (Section 2) (Figure 6.1.).

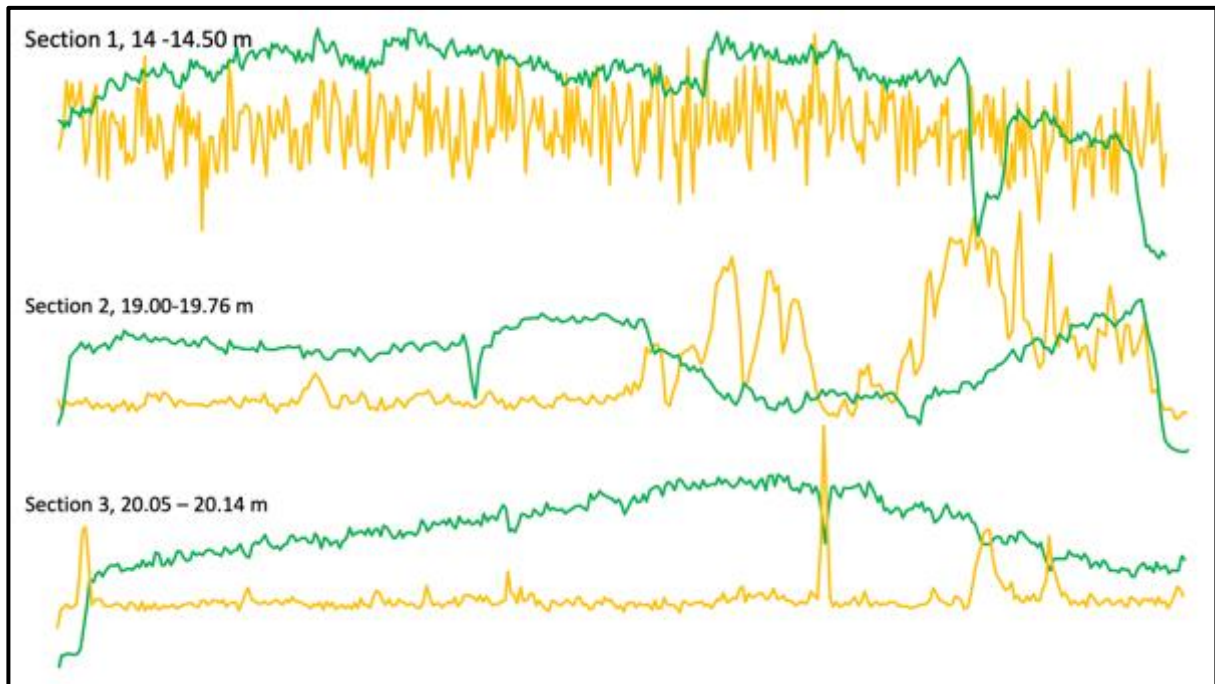


Figure 6.1. Alternating calcium (Ca = yellow) and silica (Si = green) XRF readings from sediment Sections 1, 2, and 3. Vertical scale not shown. The XRF readings are 0.5 mm at 10kV and clearly demonstrate fine resolution, alternating expressions of calcium and silica.

In addition to the simultaneously alternating peaks of Ca and Si seen in Figure 6.1., peaks in Sr/Ca are typical for aragonite layers, however the source of calcium for AZ2 sediments is not yet established for these sections (Figure 6.2.). Peaks in element ratios of Ti/Ca and Ca/(Sr+S) are proxies for increases in terrestrial detrital carbonates and silicas, respectively, washing in from watershed (Neugebauer et al., 2014). Chlorine is a proxy for salts and very positively correlates with the ratio of S/Ca which can indicate gypsum, again the exact type of calcium is not yet determined. There are significant decreases in elements at 185 mm, except the ratio Ca/(Sr+S) which peaks here. This is right before a crack in the core, however, the analyst feels these are real decreases and not a result of the crack in the section

(personal communication). Notable is the consistency of the rapid fluctuations over depth, the narrow range of their amplitude, and the tight correlations, positive and negative, of elements. While several of the element ratios contain Ca, other elements such as Sr, S, Cl, and Mn/Fe also demonstrate similar patterns providing a more robust interpretation. Mn/Fe represent terrestrial deposition and correlates negatively to Ca. High amplitude in Mn can reflect deep water turnover and can therefore be an indicator of water depth (Naeher et al., 2013; Zuo et al., 2020).

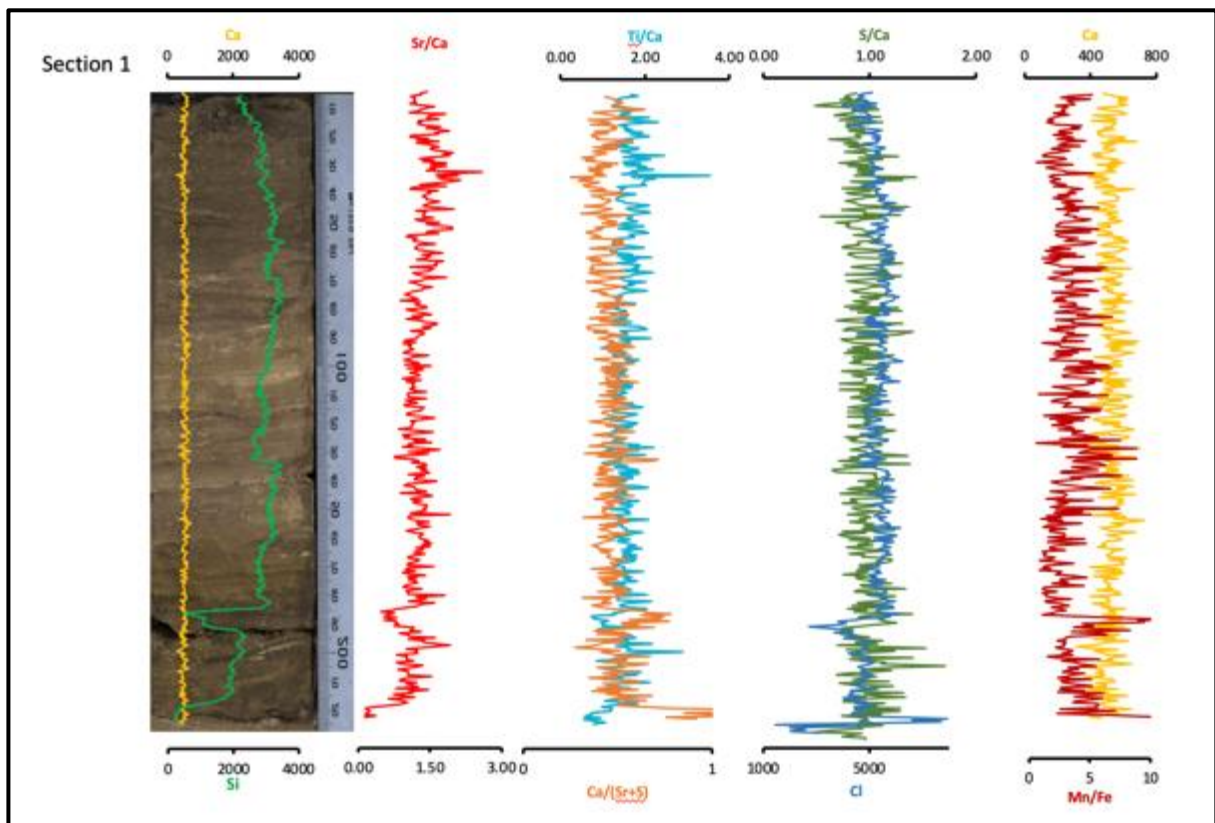


Figure 6.2. Section 1 core photo with XRF elements Ca and Si and ratios of Sr/Ca, Ti/Ca, Ca/(Sr+S), S/Ca and Cl, and Mn/FE and Ca. These are proxies for changes in depositional inputs and potential indicators of repeated seasonal cycles.

Section 2

In Section 2 the pattern of simultaneously alternating fluctuations in Ca and Si continues similar to Section 1 for the first 80 mm. The very short period and low amplitude pattern is also similar. Below 80 mm there is a dramatic shift in both period pattern and

amplitude. Silica decreases slowly, stays low then increases slowly. This could represent a decrease in available water from the watershed or a decrease in diatom productivity. Both are potentially related. During this period calcium has two dramatic episodes of multiple large spikes separated by a decrease. This is a clear divergence from the previous pattern and indicative of a significant shift in longer term sources of freshwater and higher evaporation. Strontium (Sr) was not recorded for this section, but Ti/Ca is contrasted with sulfur (S). Lake sulfur cycles are complicated with biogenic origins in microbial production, or abiotic influx from watershed erosion. In Section 2 sulfur tracks the terrestrial detrital carbonate influx represented by Ti/Ca moderately well including the shift below 80 mm, mirroring the significant drop in Ti/Ca and gradual increase, with one exception. In the middle of the decline at 107 to 110 mm, Ti/Ca has a significant peak returning to previous high level. This may indicate a strong influx of terrestrial carbon that is not reflected in the Si record. The peaks in sulfur signal may reflect reduced organic sulfides, and a decrease in sulfur signal indicate a decrease in microbial activity concurrent with a decrease in diatoms. Some portion of the sulfur signal most likely also reflects oxidized sulfur species. Further examination of sulfur species by isotope analysis or X-ray absorption near edge structure spectroscopy (XANES) is recommended (Bostick et al., 2005). Chlorine is a proxy for salts and continues to exhibit a very similar pattern to Section 1. It also continues to very positively correlate with the ratio of S/Ca. Chlorine maintains a very consistent pattern throughout Section 2 as does not reflect the major perturbations seen in Ca or S. Lastly the ratio of Mn/Fe continues to correlate negatively to Ca with high amplitude variations. After 80 mm the amplitude of Mn/Fe remains high, but the trend changes following the overall increase in calcium without the decreasing and continuing to correlate negatively. All major and most of the trace elements reflect a very similar pattern to Section 1 until 80 mm, then depict a major excursion from this pattern indicating a change in climate impacting the watershed.

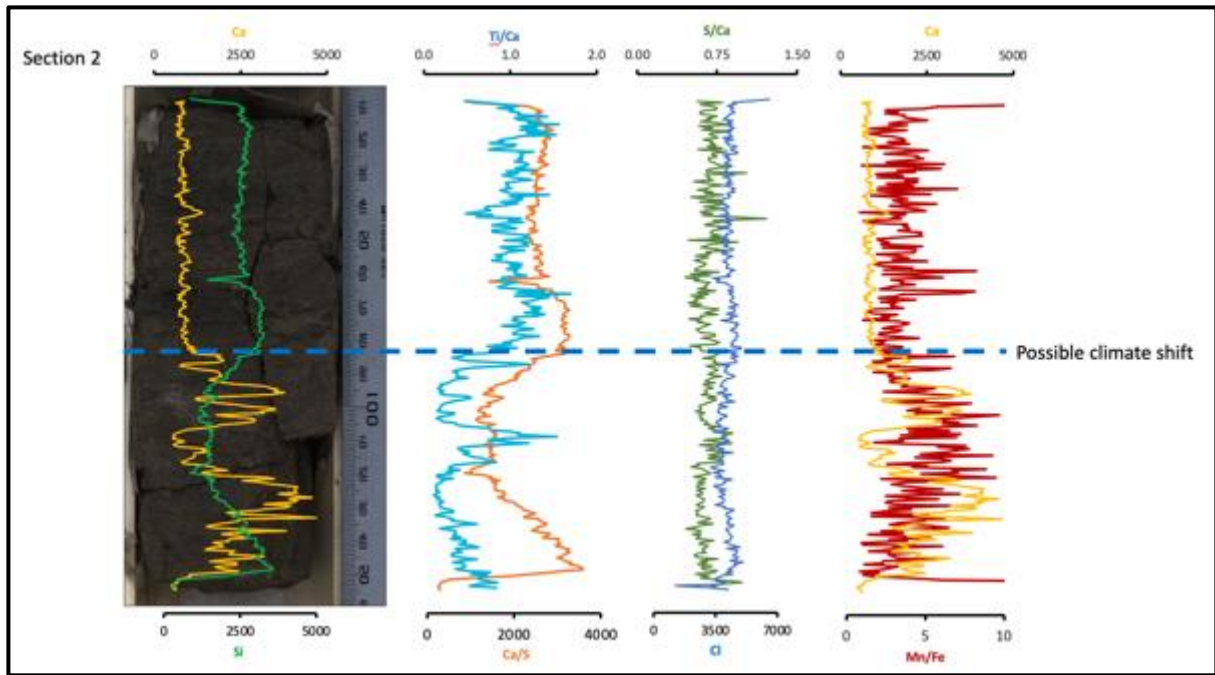


Figure 6.3. Section 2 core photo with XRF elements Ca and Si and ratios of Ti/Ca, Ca/S, S/Ca and Cl, and Mn/Fe and Ca. These are proxies for changes in depositional inputs and potential indicators of repeated seasonal cycles. A notable shift in pattern and amplitude just below 80 mm indicates a significant and abrupt change in large scale basin inputs.

Section 3

Section 3 has long continuous visible laminae of alternating light and dark colors. The XRF results are similar to Section 1 with simultaneously alternating peaks of Ca and Si throughout. The long term trend of Si increases throughout this section despite its fluctuations and amplitude remaining the same. This may reflect an overall increase in water to the lake. However, Ca has four major spikes. The largest spike occurs at 122 mm and is several orders of magnitude in scale. The other elements independent of Ca also peaking at these same four locations and similar scale are chlorine and sulfur. Iron (Fe) and silica (Si) also decrease significantly at the 122 mm depth, perhaps reflecting a significant short term arid event. The pattern and relationships of Ti/Ca and Ca/S, S/Ca and Cl, and Mn/Fe and Ca remain very similar to Section 1 indicating a stable, but seasonal climate regime dominating sediment deposition in the Al-Azraq lake.

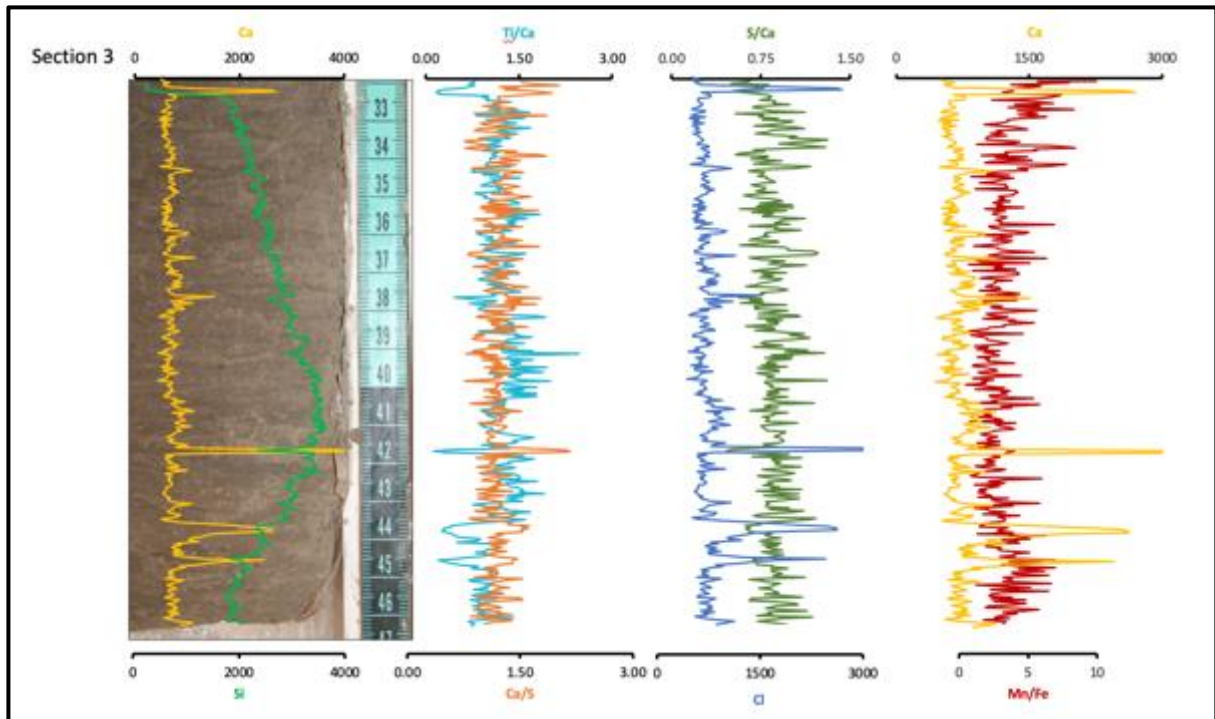


Figure 6.4. Section 3 core photo with XRF elements Ca and Si and ratios of Ti/Ca, Ca/S, S/Ca and Cl, and Mn/Fe and Ca. These are proxies for changes in depositional inputs and potential indicators of repeated seasonal cycles. A notable shift in pattern and amplitude just below 80 mm indicates a significant and abrupt change in large scale basin inputs.

Diatoms

Diatoms (Bacillariophyceae) are unicellular algae with well over a 100,000 species. They are ecologically sensitive, but can live in almost any environment that has or has had water. (Dixit et al., 1992), Their siliceous frustules are species-specific making them excellent proxies of environmental change (Battarbee, 1984). There are many factors affecting diatom preservation including: dissolution, pH, salt concentration, bacterial concentration, population, and water temperature (Lewin 1961).

SEM and EDS investigations of the AZ2 sediments sections reveal two species of centric diatoms, *Stephanodiscus* sp. and *Aulacoseira lirata* (Figure 5.10). *Stephanodiscus* is a polymorphic and one of the most diverse genera of phytoplankton. It's variability is dependent on valve size and/or ecological conditions (Theriot, Stoermer, 1981), Some *Stephanodiscus* under different ecological condition can and do exist either singly or

colonially with only a little variation in valve morphology (Stoermer et al., 1999). Because of their ability to adapt and change with conditions (Julius et al., 2010). This allows for past trophic conditions to be inferred by the different levels of available dissolved silica from the morphology of specimens preserved in sediments (Stoermer et al., 1999). *Stephanodiscus* can be found in fresh water and thrives in thermally stratified lakes. It can adapt to be successful with low Si, but requires both high phosphorous and high light conditions. It is the dominant species in winter and spring when increased turbulence suspends these relatively heavy diatoms. In Qarun Lake, Egypt domination of *Stephanodiscus* is indicative of high lake levels and increased nutrient load from freshwater inputs from Nile River flooding (Marks et al. 2017).

Aulacoseira is a planktonic freshwater diatom species found in most waters, but thrive in inland waters like lakes and ponds or rivers (Edlund *et al.*, 2008). They require high amounts of sunlight. An alkaliphilous species, or tolerant of alkaline waters. *Aulacoseira* is common in eutrophic lakes of higher temperatures and associated with warmer climates (Ehrlich, 1973; Stoermer et al., 1975; Camburn et al., 1986). It has high growth rates and therefore high demand for available silica (Kilham & Kilham 1971). They are suggestive of stable conditions, common in spring and dominant in summer. The occurrence of *Aulacoseira* with *Stephanodiscus* in low abundance indicates warm eutrophic conditions of freshwater with slightly higher salinity and alkalinity (Marks et al. 2017).

The extensive literature suggests the diatom species found in AZ2 sediments, *Stephanodiscus* and *Aulacoseira*, are subject to a range of conditions largely associated with seasonal changes in temperature and freshwater inputs. The Qarun Lake study is potentially analogous to the al-Azraq and indicates each taxa individually and in mixed conditions are indicative of different seasons (Marks et al., 2017). A detailed diatom study is required to further differentiate laminae by dominant diatom specie.

The condition of the diatoms is also informative. Some laminae consist almost entirely of diatoms with a majority of the matrix fragmented frustules. Valves appear to be broken, not dissolved. However, in Section 1, 2, and 3 there is extensive indication of post depositional halite encrustation which may be responsible for the degree of fragmentation (Figure 5.11). Further analysis into the post-depositional conditions is necessary.

Implications for Laminated Sediment

Multiple lines of evidence including: visual laminations of alternating color, grain size analysis, micro-scale fluctuations in geochemistry between wet and dry episodes, and the presence of two species of centric diatom in great quantity, which are seasonally different, all support the determination of the Al-Azraq laminations as seasonal rhythmites. The light colored laminated are composed of silt, diatoms, and organic material from the warmer spring-summer season and a darker layer composed of clay and terrestrial silica from the cooler winter season ranging in thickness from 0.5 to 1.5 cm. Whether these coupled deposits can be defined as annual varves is not yet established as the specific species of calcium and sulfur are yet to be determined. Determining if all laminations are seasonal rhythmites is also not definitive and will be further investigated with micro-facies analysis. It is reasonable to presume there is a great deal of variation in depositional sequences and timing throughout the evolution of this paleolake (Ahmad, Davies, 2021). Age control is also necessary at the micro-scale to support the seasonal determination and to better place these sediments into the wider paleoenvironmental context of the Dead Sea basin laminated sediments and beyond.

Implications of Paleoenvironmental Variability and Regional Implications

The lacustrine sediments of the Al-Azraq basin demonstrate a wide range of paleoenvironments influenced basin evolution and shift in climate regime (Ahmad, Davies 2017, 2021). The laminated sediments investigated here contribute additional understanding

to the paleoenvironmental history of the basin and broader region. The rhythmites indicate both long-term stability of the paleoenvironment, and also signal abrupt changes in climate impacting the watershed (Kemp, 2003). The long-term stability of the regional climate regime is reflected in the meters of laminated sediment, of which only three short sections were examined in detail here. This is supported by the long record of steady alternating fluctuations of Ca and Si, as well as the consistency in the micro-scale short period and consistent low amplitude. The presence of meters of diatomite in the Al-Azraq first reported by Al'ali (2009), is expressed in 11 meters of AZ2 lacustrine sediments.

There is also evidence of abrupt changes in temperature and precipitation regime. The range of sediment grain size throughout the core while overall falling within the sandy silt range, show horizons of clays, aeolian, phases, and episodes of high water levels and drying episodes. Within the three sections examined here a significant potential climate shift is recorded in Section 2 below 80 mm with abrupt change of pattern and shift to different thresholds. Additional evidence of changes in climate is revealed by the EDS analysis of diatom samples showing extensive halite encrusting diatom frustules post-deposition. This indicates a period of high temperature. It also has implication for other carbonates and the sequence of evaporation.

Wet winters and dry summers characterize the climate of the eastern Mediterranean basin. The Jordan Plateau and eastern Mediterranean regional falls within a climatic transition zone between the fluctuations of the Northern jet Stream (NJS) and Subtropical Jet Stream (STJ). Throughout the Paleocene these circulation patterns have shifted in dominance producing more or less moisture between glaciations and non-glacial periods. The Dead Sea and its precursor lakes have long been viewed as one of the best gauges of precipitation history for the region (Neev et al., 1967, Migowsk, 2006). A study by Migowsk indicate two major wet phases along with a number of abrupt arid phases in the Jordan Valley and Dead

Sea region. The recent work by Ben Dor (2019) continue to demonstrate the regional import of laminated sediment. While the work by Lopez et al. (2016) and responses by Bookman (2020) stress the need to examine the basis for determining coupled laminae as seasonal deposition. It is reasonable to assume not all rhythmites represent strictly annual deposition, but rather seasonal being a wet and dry season. It is known from the present day Al-Azraq basing there are some years without any precipitation, and research has noted several very wet phases in the Dead Sea where water levels during the Holocene rose above 400 m bmsl, Neev and Emery (1967). It should be pointed out that the Dead Sea is a very large and very deep basin where it is unlikely annual depositional layers will cover the entire basin.

One line of critic of the annual designation of Dead Sea varves relies on demonstrating the source of pollen is not seasonal, but washed in and mixed from the surrounding watershed or aeolian in origin. However, in the Al-Azraq basin biological evidence for seasonality comes from *in situ* diatom production which is tied directly to lake environmental conditions (Gasse, 1994; Gasse et al., 1983; Gasse, et al. 1997). The diatom species found in the AZ2 laminations, *Stephanodiscus sp.* and *Aulacoseira lirata* are deposited in late summer earl winter, after the water settles down, the elongated centric diatoms tend to be spring and early summer, respectfully. The formation of a matrix bed of diatom frustules (Si, O) in matrix of crushed frustules, Fe-Mg-Al clay, and encrusted with halite (NaCl) indicate arid conditions occurring post-deposition.

From a smaller playa basin, Lake Qarun in the Feiyum Oasis west of the Nile River in Egypt perhaps more similar to the Al-Azraq, comes an additional study of note f (Marks et al. 2016; Zalat et al., 2017). The laminated sediments from this sediment core are also dominated by *Aulacoseira spp.* and *Stephanodiscus spp.* assemblages, but *Cyclostephanos dubius* and *Cyclotella meneghiniana* are also common. Two main factors in diatom

preservation at Quran are lake levels and salt levels. These would suggest similar climate conditions across the transitional region of the eastern Mediterranean (Flower, 2009).

7. CONCLUSIONS

Conclusions

The results of this investigation demonstrate the presence of laminated sediments and through robust examination by multiple analyses determine them to be seasonal rhythmites. These sediments reflect significant changes in depositional processes, moisture fluxes, and temperature regimes. These records document recurring processes and abrupt changes in the climate of the Quaternary evolution of this ancient saline lake. The sediment and mineralogical analysis of the AZ2 core sediments indicate changes in energy, environment, episodes of tectonic activity, terrestrial and aeolian inputs, depositional changes, and high evaporation phases. AZ2 core records approximately 11 meters with laminated sediments having alternating laminae of light and dark deposits composed of clays, silts, and diatoms. The detailed analyses of three sections of laminates provides a window into fine resolution lake processes.

Arid lake sediments are a complex mix of autochthonous biotic (microbial reduction) and abiotic (reduction and precipitation) processes and allochthonous inputs from the surrounding watershed which bring terrestrial minerals and organic detritus into the basin (Gierlowski-Kordesch, 2010). Calcium and silica are major components within the AZ2 sediments and rhythmites. Calcium is high throughout catchment (basalt and carbonate sources) and from aeolian sources. Is it also produced through evaporation within the lake. Silica also has multiple origins from erosion in the watershed and the biogenic productivity of diatoms. Mapping the rhythmite elements through XRF indicate variations are dominated by the seasonal deposition of additional terrigenous material, as seen primarily through Al, Fe, K, Sr and Ti. The Ca, Cl, S and to some extent a portion of Si are indicative of internal lake processes.

This research identifies the presence of laminae in AZ2 sediments and further demonstrates these laminated sediments are seasonal rhythmites. The rhythmites are composed of alternating calcium and silica. There are other signals in the element data requiring further investigation to identify the types of carbonates (calcite, aragonite, halite, gypsum, etc.), which will provide additional data on the seasonality signal. Additionally the presence of two species of differing seasonal express also supports seasonal differences in depositional timing. This research demonstrates, contrary to Ben Dor (2016), the Dead Sea is not the only deep hypersaline lake in the region containing finely laminated rhythmites and the potential for a high resolution paleoenvironmental record. This research also suggests Qarun Lake, Egypt may also serve as an analog for the Al-Azraq paleolake sharing similar diatom sequences.

APPENDIX A

Grain Size Data

AZ2 Grain Size Raw Data						
Depth	Tot Clay	Tot silt	tot sand	Tot	Sheppard	Folk
5.51	10.44	64.77	24.78	100.00	SnSilt	SnSilt
5.55	8.43	68.90	22.66	99.99	SnSilt	SnSilt
5.59	9.16	66.81	24.03	100.00	SnSilt	SnSilt
5.64	11.28	64.56	24.16	100.00	SnSilt	SnSilt
5.68	7.83	60.72	31.45	99.99	SnSilt	SnSilt
5.70	8.15	61.31	30.55	100.01	SnSilt	SnSilt
5.72	14.57	64.22	21.21	100.00	SnSilt	SnSilt
5.79	11.16	60.38	28.47	100.00	SnSilt	SnSilt
5.83	9.33	64.15	26.52	100.00	SnSilt	SnSilt
5.89	11.40	64.42	24.18	100.00	SnSilt	SnSilt
5.94	9.39	54.56	36.04	100.00	SnSilt	SnSilt
6.00	12.76	65.63	21.63	100.02	SnSilt	SnSilt
6.05	12.39	69.20	18.41	100.00	SnSilt	SnSilt
6.09	11.94	57.40	30.66	100.00	SnSilt	SnSilt
6.51	9.61	59.58	30.80	100.00	SnSilt	SnSilt
6.56	11.58	60.99	27.44	100.01	SnSilt	SnSilt
6.61	10.03	65.38	24.59	99.99	SnSilt	SnSilt
6.68	11.58	63.32	25.09	99.99	SnSilt	SnSilt
6.75	11.79	51.01	37.21	100.01	SnSilt	SnSilt
6.82	19.81	64.70	15.51	100.02	ClayeySilt	SnSilt
6.88	17.87	57.04	25.10	100.02	SnSilt	SnSilt
6.94	15.36	60.90	23.75	100.01	SnSilt	SnSilt
6.99	20.50	66.81	12.69	100.00	ClayeySilt	SnSilt
7.00	16.60	59.39	24.00	99.99	SnSilt	SnSilt
7.10	9.61	59.58	30.80	100.00	SnSilt	SnSilt
7.15	9.61	59.58	30.80	100.00	SnSilt	SnSilt
7.30	9.61	59.58	30.80	100.00	SnSilt	SnSilt
7.34	9.61	59.58	30.80	100.00	SnSilt	SnSilt
7.39	9.61	59.58	30.80	100.00	SnSilt	SnSilt
7.42	9.61	59.58	30.80	100.00	SnSilt	SnSilt
7.52	10.74	60.73	28.52	99.99	SnSilt	SnSilt
7.57	8.55	64.18	27.27	100.00	SnSilt	SnSilt
7.64	12.06	62.38	25.55	99.99	SnSilt	SnSilt
7.68	10.35	61.51	28.14	100.00	SnSilt	SnSilt
8.02	18.52	71.87	9.60	99.99	ClayeySilt	Silt
8.09	10.45	62.16	27.38	99.99	SnSilt	SnSilt
8.22	14.16	66.95	18.88	99.99	SnSilt	SnSilt
8.26	13.97	67.76	18.27	100.00	SnSilt	SnSilt
8.51	10.53	62.98	26.47	99.99	SnSilt	SnSilt
8.56	40.04	49.23	10.74	100.01	SnSilt	SnSilt
8.67	10.34	59.23	30.42	100.00	SnSilt	SnSilt
8.70	15.60	67.48	16.92	100.00	SnSilt	SnSilt
8.74	14.15	68.40	17.45	99.99	SnSilt	SnSilt
8.90	14.97	65.36	19.67	100.00	SnSilt	SnSilt

9.09	15.75	60.31	23.94	100.00	SnSilt	SnSilt
9.83	13.00	58.51	28.48	99.99	SnSilt	SnSilt
9.89	9.72	55.82	34.46	99.99	SnSilt	SnSilt
9.92	8.99	51.98	39.04	100.01	SnSilt	SnSilt
9.97	8.59	65.76	25.66	100.01	SnSilt	SnSilt
10.06	3.26	59.12	37.63	100.01	SnSilt	SnSilt
10.10	3.12	50.87	46.01	100.00	SnSilt	SnSilt
10.18	3.11	53.48	43.40	100.00	SnSilt	SnSilt
10.22	2.00	49.19	48.82	100.01	SnSilt	SnSilt
10.26	2.52	52.12	45.39	100.02	SnSilt	SnSilt
10.32	6.54	62.99	30.47	100.00	SnSilt	SnSilt
10.38	3.03	62.06	34.90	100.00	SnSilt	SnSilt
10.47	6.85	63.32	29.85	100.01	SnSilt	SnSilt
10.53	9.09	71.27	19.64	100.00	SnSilt	SnSilt
11.00	2.03	62.95	35.02	100.00	SnSilt	SnSilt
11.04	1.77	67.32	30.92	100.01	SnSilt	SnSilt
11.10	1.62	56.31	42.07	100.00	SnSilt	SnSilt
11.20	1.06	58.77	40.17	100.00	SnSilt	SnSilt
11.25	0.57	59.52	39.90	100.00	SnSilt	SnSilt
11.32	1.03	52.89	46.09	100.01	SnSilt	SnSilt
11.42	2.52	67.46	30.02	100.00	SnSilt	SnSilt
11.49	0.87	49.04	50.09	99.99	SiltySand	SiltySand
12.60	4.31	61.52	34.19	100.02	SnSilt	SnSilt
12.65	3.11	68.55	28.34	100.00	SnSilt	SnSilt
12.67	2.76	57.44	39.81	100.00	SnSilt	SnSilt
12.69	3.55	54.99	41.47	100.01	SnSilt	SnSilt
12.75	4.08	60.38	35.54	100.00	SnSilt	SnSilt
12.78	8.46	54.82	36.74	100.01	SnSilt	SnSilt
12.86	7.92	57.86	34.22	100.00	SnSilt	SnSilt
12.88	8.46	54.82	36.74	100.01	SnSilt	SnSilt
12.89	8.46	54.82	36.74	100.01	SnSilt	SnSilt
13.01	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.01	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.04	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.10	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.11	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.16	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.23	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.31	6.31	50.85	42.83	99.99	SnSilt	SnSilt
14.38	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.46	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.56	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.63	8.46	54.82	36.74	100.01	SnSilt	SnSilt
14.68	8.46	54.82	36.74	100.01	SnSilt	SnSilt
15.77	6.24	83.57	10.20	100.00	Silt	SnSilt
15.84	5.51	72.69	21.80	100.00	SnSilt	SnSilt

15.86	7.27	67.25	25.48	100.00	SnSilt	SnSilt
17.03	10.66	67.55	21.80	100.01	SnSilt	SnSilt
17.08	5.85	66.41	27.76	100.01	SnSilt	SnSilt
17.12	6.18	65.93	27.84	99.95	SnSilt	SnSilt
17.18	7.38	62.11	30.29	99.78	SnSilt	SnSilt
17.21	10.15	62.96	26.89	100.00	SnSilt	SnSilt
17.30	13.08	68.44	18.44	99.96	SnSilt	SnSilt
17.38	7.36	65.66	26.98	100.00	SnSilt	SnSilt
17.42	9.46	67.17	23.37	99.99	SnSilt	SnSilt
17.45	5.52	68.09	26.38	99.99	SnSilt	SnSilt
17.50	7.62	70.75	21.65	100.02	SnSilt	SnSilt
17.59	9.22	69.15	21.64	100.01	SnSilt	SnSilt
17.66	9.92	76.34	13.73	99.99	Silt	SnSilt
17.71	6.62	57.68	35.71	100.01	SnSilt	SnSilt
17.76	8.53	69.36	22.10	99.99	SnSilt	SnSilt
17.81	5.78	61.55	32.67	100.00	SnSilt	SnSilt
17.85	7.55	74.21	18.23	99.99	SnSilt	SnSilt
17.94	5.90	53.98	40.12	100.00	SnSilt	SnSilt
18.00	7.43	73.64	18.92	99.99	SnSilt	SnSilt
18.06	8.69	56.96	34.38	100.03	SnSilt	SnSilt
18.14	10.69	58.51	30.83	100.03	SnSilt	SnSilt
18.20	8.66	49.29	42.04	99.99	SnSilt	SnSilt
19.02	7.71	65.70	26.60	100.01	SnSilt	SnSilt
19.09	6.58	64.77	28.66	100.01	SnSilt	SnSilt
19.16	9.32	73.90	16.77	99.99	SnSilt	SnSilt
19.20	7.24	54.63	38.12	99.99	SnSilt	SnSilt
19.23	9.57	67.07	23.36	100.00	SnSilt	SnSilt
19.30	6.50	60.38	33.66	100.54	SnSilt	SnSilt
19.33	10.37	74.14	15.50	100.01	SnSilt	SnSilt
19.41	8.45	57.28	34.26	99.99	SnSilt	SnSilt
20.50	11.56	67.53	20.91	100.00	SnSilt	SnSilt
20.57	14.24	66.00	19.77	100.01	SnSilt	SnSilt
20.61	12.67	64.35	22.97	100.00	SnSilt	SnSilt
22.00	14.41	72.27	13.33	100.00	ClayeySilt	SnSilt
22.04	14.12	71.94	13.93	99.99	ClayeySilt	SnSilt
22.05	15.20	67.72	17.08	100.00	SnSilt	SnSilt
22.20	12.82	68.86	18.33	100.00	SnSilt	SnSilt
22.28	16.07	74.64	9.29	100.00	ClayeySilt	Silt
22.38	11.50	50.89	37.59	99.98	SnSilt	SnSilt
22.44	17.01	69.81	13.17	99.99	SnSilt	SnSilt
22.54	11.55	52.07	36.38	100.00	SnSilt	SnSilt
22.56	16.55	62.27	21.18	100.00	SnSilt	SnSilt
22.62	9.75	55.75	34.50	100.00	SnSilt	SnSilt
22.64	12.97	56.88	30.14	100.00	SnSilt	SnSilt
22.92	8.87	46.95	44.58	100.40	SnSilt	SnSilt
23.03	11.05	55.83	33.54	100.41	SnSilt	SnSilt

23.05	10.60	72.08	17.70	100.38	SnSilt	SnSilt
23.15	11.54	43.03	46.02	100.59	SnSilt	SnSilt
23.29	7.22	51.99	18.93	78.14	SnSilt	SnSilt
23.34	7.59	64.43	28.32	100.34	SnSilt	SnSilt
23.56	11.55	76.70	11.72	99.97	Silt	SnSilt
23.59	11.55	76.70	11.72	99.97	Silt	SnSilt
23.62	7.87	72.54	19.56	99.97	SnSilt	SnSilt
23.66	8.27	70.39	21.33	99.99	SnSilt	SnSilt
23.69	6.72	39.78	53.47	99.98	SiltySand	SiltySand
23.73	11.61	63.01	25.35	99.97	SnSilt	SnSilt
23.76	10.33	69.29	20.36	99.98	SnSilt	SnSilt
23.78	16.30	66.09	17.57	99.96	SnSilt	SnSilt
23.81	16.93	53.34	29.68	99.95	SnSilt	SnSilt
25.24	11.32	85.33	3.35	100.00	Silt	Silt
25.28	13.75	81.95	4.30	100.00	Silt	Silt
25.29	10.49	80.17	9.34	100.00	Silt	Silt
25.32	16.39	83.49	0.11	99.99	Silt	Silt
25.34	18.23	79.42	2.34	100.00	Silt	Silt
25.38	15.65	80.67	3.68	100.00	Silt	Silt
25.41	14.21	82.40	3.39	100.00	Silt	Silt
25.43	16.13	80.07	3.80	100.00	Silt	Silt
25.44	12.20	73.66	14.13	99.99	SnSilt	SnSilt
25.45	13.00	78.85	8.15	100.00	SnSilt	SnSilt
25.46	11.76	83.24	4.99	99.99	Silt	Silt
25.48	12.63	79.03	8.34	100.01	Silt	Silt
25.50	11.66	56.10	32.23	99.99	SnSilt	SnSilt
25.51	13.20	82.93	3.87	100.00	Silt	Silt
25.55	13.09	77.94	8.95	99.99	Silt	Silt
25.92	11.76	70.79	17.46	100.00	SnSilt	SnSilt
35.24	18.43	76.60	4.97	99.99	Silt	Silt
35.34	14.60	72.56	12.83	100.00	ClayeySilt	SnSilt
35.45	11.92	73.32	14.75	100.00	SnSilt	SnSilt
35.54	12.47	71.65	15.88	99.99	SnSilt	SnSilt
35.64	12.73	71.58	15.70	100.01	SnSilt	SnSilt
35.68	13.76	73.72	12.49	99.97	ClayeySilt	SnSilt
35.76	13.53	76.67	9.81	100.00	Silt	Silt
35.87	13.90	77.03	9.08	100.01	Silt	Silt
35.96	27.59	55.21	17.19	99.99	ClayeySilt	SandyMud
35.98	11.89	65.80	22.31	100.00	SnSilt	SnSilt

APPENDIX B

X-Ray Florescence Raw Data

d\VPData_Tube	10	10	10/6/15	57.5	243	30	2474	59	32	31	489	50	4153	267	840	72	841	68	949	80	583	91	495	93	6005	132	5234	313
d\VPData_Tube	10	10	10/6/15	58	144	31	2505	61	8	32	543	51	4121	264	774	71	986	68	946	79	769	91	280	92	5895	132	5107	312
d\VPData_Tube	10	10	10/6/15	58.5	168	31	2647	62	-1	32	452	51	4592	273	786	73	833	69	899	79	704	91	138	91	6078	133	4593	324
d\VPData_Tube	10	10	10/6/15	59	220	32	2588	62	4	31	609	50	4370	271	706	73	790	69	825	79	516	92	307	94	6096	134	4764	322
d\VPData_Tube	10	10	10/6/15	59.5	250	32	2611	63	-4	31	408	51	4334	275	639	73	820	68	876	80	607	91	379	93	6199	134	5037	329
d\VPData_Tube	10	10	10/6/15	60	169	31	2812	64	-6	32	509	51	4141	277	792	73	735	69	873	81	591	92	169	93	6016	133	5415	336
d\VPData_Tube	10	10	10/6/15	60.5	146	31	2461	61	14	33	372	49	4330	270	612	72	828	68	752	78	604	91	201	92	5655	131	4682	318
d\VPData_Tube	10	10	10/6/15	61	95	28	1869	54	3	30	435	49	4340	286	559	70	741	66	680	78	568	91	164	92	4788	127	5624	313
d\VPData_Tube	10	10	10/6/15	61.5	122	26	1430	48	6	29	382	48	3741	289	422	67	713	65	617	76	545	89	189	90	4365	123	4865	313
d\VPData_Tube	10	10	10/6/15	62	142	30	2170	57	7	31	532	50	3537	292	634	73	781	68	841	80	605	92	166	93	5573	132	5645	352
d\VPData_Tube	10	10	10/6/15	62.5	166	31	2537	60	34	31	537	51	4138	280	708	74	805	70	785	80	620	93	154	95	6174	133	5089	338
d\VPData_Tube	10	10	10/6/15	63	203	31	2561	61	22	31	622	51	4212	264	805	74	978	70	905	80	536	91	127	93	6022	133	4769	314
d\VPData_Tube	10	10	10/6/15	63.5	252	32	2557	61	25	31	679	51	4452	262	825	75	1039	71	888	80	635	92	517	94	6333	134	4797	310
d\VPData_Tube	10	10	10/6/15	64	261	34	2870	64	-4	31	518	51	4483	277	837	75	904	71	1012	81	581	92	491	95	6734	135	4954	334
d\VPData_Tube	10	10	10/6/15	64.5	289	34	2840	64	-3	31	596	50	4230	282	904	75	804	71	1122	81	558	91	235	93	6745	137	5408	340
d\VPData_Tube	10	10	10/6/15	65	305	33	2955	64	-15	32	500	51	4516	272	1015	75	888	70	951	81	608	92	111	94	6837	138	5265	327
d\VPData_Tube	10	10	10/6/15	65.5	197	33	3008	66	-1	32	460	50	4266	266	888	76	682	69	1158	81	579	91	133	93	6533	137	5172	320
d\VPData_Tube	10	10	10/6/15	66	252	33	2941	65	-15	32	402	50	4390	265	975	75	790	70	832	81	477	91	302	93	6863	137	5309	318
d\VPData_Tube	10	10	10/6/15	66.5	310	35	2957	66	7	34	432	51	4530	279	1057	77	680	70	988	81	698	92	145	92	6814	137	5153	334
d\VPData_Tube	10	10	10/6/15	67	252	32	2993	65	0	32	416	51	4468	279	837	76	762	70	1003	83	577	93	216	94	6903	138	5322	338
d\VPData_Tube	10	10	10/6/15	67.5	213	34	3140	67	32	32	521	51	4341	266	987	77	916	70	1141	82	637	92	204	96	7248	140	5235	324
d\VPData_Tube	10	10	10/6/15	68	348	34	3165	67	10	33	412	51	4378	267	1229	77	852	70	963	83	517	92	213	94	7425	141	5289	324
d\VPData_Tube	10	10	10/6/15	68.5	255	34	3165	68	15	35	505	51	4554	269	1163	78	834	71	1005	82	641	92	267	94	7401	140	5131	325
d\VPData_Tube	10	10	10/6/15	69	219	33	3192	67	-2	32	467	51	4762	260	960	75	936	71	1128	81	640	93	111	94	7163	138	4804	314
d\VPData_Tube	10	10	10/6/15	69.5	255	34	3173	68	7	34	488	51	4249	273	977	77	798	70	1086	82	613	93	87	94	7067	138	5515	333
d\VPData_Tube	10	10	10/6/15	70	323	35	3274	67	9	32	575	52	4599	274	952	76	745	71	929	82	476	93	125	94	7240	140	5123	335
d\VPData_Tube	10	10	10/6/15	70.5	240	33	3219	67	11	32	515	50	4546	260	873	75	714	70	954	82	635	92	222	95	6883	138	5206	315
d\VPData_Tube	10	10	10/6/15	71	270	33	3153	67	-7	33	477	52	4572	265	1107	77	751	70	872	82	595	92	155	93	6980	138	5182	318
d\VPData_Tube	10	10	10/6/15	71.5	307	34	3118	66	35	33	538	51	4097	256	1071	78	710	70	990	83	515	91	205	93	7687	140	5438	350
d\VPData_Tube	10	10	10/6/15	72	248	34	3190	67	7	32	546	51	4716	266	971	76	910	71	1186	82	746	93	75	93	7161	139	5078	321
d\VPData_Tube	10	10	10/6/15	72.5	219	34	3091	67	12	32	551	51	4346	274	1022	76	899	71	934	82	620	93	221	94	6861	138	5544	332
d\VPData_Tube	10	10	10/6/15	73	228	34	3116	67	-5	33	485	51	4423	277	863	74	993	70	1072	81	746	92	265	94	6669	137	5547	336
d\VPData_Tube	10	10	10/6/15	73.5	226	32	3143	67	-4	33	530	53	3829	279	968	75	875	70	893	81	552	91	231	93	6490	136	6456	340
d\VPData_Tube	10	10	10/6/15	74	210	35	3214	69	-22	34	490	52	4446	280	919	76	864	71	1012	81	578	91	187	94	6462	136	5457	341
d\VPData_Tube	10	10	10/6/15	74.5	288	34	3095	66	18	31	594	51	4819	273	1037	76	864	70	957	81	545	91	147	93	6463	136	5104	326
d\VPData_Tube	10	10	10/6/15	75	302	33	3096	66	-7	32	575	52	4615	263	1132	74	916	69	827	81	561	93	230	94	6562	137	5249	315
d\VPData_Tube	10	10	10/6/15	75.5	199	33	3203	68	-6	35	497	52	4693	267	851	75	818	71	940	80	514	91	214	93	6630	136	5087	320
d\VPData_Tube	10	10	10/6/15	76	303	34	3153	67	47	33	435	50	4445	275	889	76	781	71	994	81	664	92	139	93	6600	138	5588	331
d\VPData_Tube	10	10	10/6/15	76.5	222	34	3137	67	10	32	558	52	4347	284	840	75	973	71	1049	82	481	93	134	95	6678	137	5661	347
d\VPData_Tube	10	10	10/6/15	77	199	33	3173	67	20	31	523	50	4656	263	857	75	876	71	918	82	540	92	223	94	6390	137	4999	317
d\VPData_Tube	10	10	10/6/15	77.5	253	32	3270	66	46	30	582	52	4693	262	882	75	798	70	821	81	633	93	300	95	6547	136	5231	316
d\VPData_Tube	10	10	10/6/15	78	274	33	3261	67	36	32	584	52	4757	272	758	76	794	70	1061	82	709	92	100	93	6544	137	5255	328
d\VPData_Tube	10	10	10/6/15	78.5	253	34	3251	67	21	32	573	51	4672	275	1060	76	950	71	1093	82	592	93	164	95	6723	138	5273	335
d\VPData_Tube	10	10	10/6/15	79	206	33	3166	67	39	33	623	52	4648	267	805	75	919	71	994	81	642	93	158	94	7557	142	4800	323
d\VPData_Tube	10	10	10/6/15	79.5	194	32	3128	66	3	32	640	53	4590	265	956	75	867	70	882	81	568	92	146	94	7792	142	5119	320
d\VPData_Tube	10	10	10/6/15	80	242	32	3097	66	18	34	759	54	4396	264	827	75	979	71	892	81	585	93	294	94	8274	142	5194	325
d\VPData_Tube	10	10	10/6/15	80.5	221	34	3035	66	2	33	695	53	4466	262	932	75	964	70	876	81	619	92	239	93	7601	140	5236	315
d\VPData_Tube	10	10	10/6/15	81	179	34	3157	68	-23	34	570	53	4376	278	1074	75	1160	72	853	81	618	92	106	93	6980	139	5632	339
d\VPData_Tube	10	10	10/6/15	81.5	246	33	3010	65	34	32	582	52	4599	265	783	74	909	71	1044	80	775	91	232	90	6716	136	4971	314
d\VPData_Tube	10	10	10/6/15	82	217	32	2954	66	-17	34	729	54	4923	272	905	75	1226	73	972	80	652	91	151	92	6413	136	4759	325
d\VPData_Tube	10	10	10/6/15	82.5	227	34	3143	68	-7	34	647	54	3878	281	925	75	1075	72	970	81	550	92	152	93	6635	137	6023	346
d\VPData_Tube	10	10	10/6/15	83	266	32	3158	66	-27	32	1000	57	4695	260	812	76	1300	76	1075	81	754	91	449	92	6566	135	5065	316
d\VPData_Tube	10	10	10/6/15	83.5	286	33	2921	65	-19	33	1444	61	4408	2														

d\WFData_Tube	10	10	10/6/15	107.5	76	28	1595	51	-5	31	497	51	3975	316	395	70	589	66	641	80	692	94	118	93	3535	120	4968	356
d\WFData_Tube	10	10	10/6/15	108	107	28	1495	50	16	31	442	50	3839	310	387	70	530	66	815	78	595	91	128	92	3570	120	5170	338
d\WFData_Tube	10	10	10/6/15	108.5	90	26	1526	50	-6	29	584	51	3858	315	379	69	719	67	735	79	600	92	63	92	3672	121	5258	354
d\WFData_Tube	10	10	10/6/15	109	75	25	1553	50	33	30	546	50	4276	293	403	68	869	67	644	78	591	90	127	92	3385	118	4748	319
d\WFData_Tube	10	10	10/6/15	109.5	93	29	1430	50	0	31	479	51	3984	346	415	69	630	66	621	78	639	91	261	91	3442	117	5333	377
d\WFData_Tube	10	10	10/6/15	110	73	26	1323	48	9	31	487	50	4034	319	362	68	544	65	660	79	529	91	208	92	3596	120	5129	342
d\WFData_Tube	10	10	10/6/15	110.5	75	26	1523	49	1	30	421	51	3648	318	531	69	722	66	592	78	547	91	279	92	3595	120	5619	356
d\WFData_Tube	10	10	10/6/15	111	94	27	1415	49	-23	32	1040	57	3427	287	239	70	1471	75	741	78	609	91	252	93	3578	119	5075	328
d\WFData_Tube	10	10	10/6/15	111.5	123	28	1484	49	14	30	1016	57	3957	297	260	69	1373	74	543	78	578	91	236	91	3786	120	5036	335
d\WFData_Tube	10	10	10/6/15	112	145	28	1456	50	-28	32	1030	57	3610	296	266	69	1585	75	685	77	575	90	169	91	3629	120	5366	338
d\WFData_Tube	10	10	10/6/15	112.5	90	28	1435	49	15	32	946	55	4485	290	361	70	1163	72	612	77	615	91	355	93	3839	120	4338	319
d\WFData_Tube	10	10	10/6/15	113	82	27	1567	51	-17	31	888	55	3840	294	419	69	1146	71	519	78	613	91	164	91	3771	120	5077	333
d\WFData_Tube	10	10	10/6/15	113.5	130	28	1424	50	-2	33	795	54	3509	302	339	67	1182	70	598	78	521	90	208	92	3522	120	4898	330
d\WFData_Tube	10	10	10/6/15	114	105	27	1498	49	10	30	796	54	3956	315	280	68	1018	70	569	79	589	91	228	92	3440	119	5335	352
d\WFData_Tube	10	10	10/6/15	114.5	117	28	1536	50	-5	31	618	53	3695	323	230	68	927	69	631	80	620	93	225	93	3704	121	5421	367
d\WFData_Tube	10	10	10/6/15	115	138	27	1543	49	4	29	609	52	3873	287	368	68	805	67	588	78	581	90	84	91	3305	118	4875	314
d\WFData_Tube	10	10	10/6/15	115.5	124	28	1440	49	19	31	698	53	3917	330	303	69	859	69	721	79	563	91	186	92	3634	121	5256	366
d\WFData_Tube	10	10	10/6/15	116	96	26	1331	48	-4	31	770	55	3930	330	423	68	1375	72	749	78	530	91	136	93	3250	118	5373	364
d\WFData_Tube	10	10	10/6/15	116.5	60	25	1210	45	6	30	976	57	3553	288	298	67	1681	74	570	78	620	91	291	92	3281	116	5355	320
d\WFData_Tube	10	10	10/6/15	117	38	24	1031	44	-26	31	922	56	3225	293	51	66	1812	76	555	76	684	91	199	92	2966	114	5364	318
d\WFData_Tube	10	10	10/6/15	117.5	15	25	1022	43	28	29	1244	59	3618	307	166	67	2108	78	600	77	661	91	189	91	2644	112	5235	343
d\WFData_Tube	10	10	10/6/15	118	72	24	1016	42	44	30	876	55	3982	325	41	66	1373	73	574	77	766	93	145	93	3088	115	4843	348
d\WFData_Tube	10	10	10/6/15	118.5	38	24	858	41	16	29	980	57	3359	336	23	66	1503	74	596	75	531	90	166	91	2380	109	5435	360
d\WFData_Tube	10	10	10/6/15	119	91	27	1291	47	1	30	1183	59	3702	298	299	68	1718	76	555	78	646	91	308	91	3246	117	5190	337
d\WFData_Tube	10	10	10/6/15	119.5	45	27	1485	51	-2	33	1951	66	3834	254	220	68	3255	88	489	78	587	90	76	91	3462	117	4835	302
d\WFData_Tube	10	10	10/6/15	120	48	28	1424	51	-27	34	1845	66	3495	261	342	67	3477	89	528	77	553	91	234	94	3511	119	5348	312
d\WFData_Tube	10	10	10/6/15	120.5	98	29	1673	52	-16	32	1653	63	4425	266	326	69	2551	82	861	78	599	91	148	91	3929	120	4462	311
d\WFData_Tube	10	10	10/6/15	121	70	29	1484	51	3	33	1914	66	3945	259	335	69	3032	86	589	77	504	90	105	91	3674	119	4910	305
d\WFData_Tube	10	10	10/6/15	121.5	106	28	1592	51	-16	33	1873	66	4041	259	356	69	3317	88	718	78	638	91	151	91	3998	121	4803	308
d\WFData_Tube	10	10	10/6/15	122	115	29	1550	52	-37	34	2359	70	3901	259	196	69	3718	93	611	78	524	89	163	90	3582	118	5034	313
d\WFData_Tube	10	10	10/6/15	122.5	91	28	1557	52	-27	34	2369	71	3415	259	219	70	4162	95	554	78	571	90	226	92	3687	118	5360	320
d\WFData_Tube	10	10	10/6/15	123	83	29	1579	52	-32	34	2290	70	3624	248	213	70	4069	94	591	77	530	90	130	91	3706	119	5020	302
d\WFData_Tube	10	10	10/6/15	123.5	78	28	1555	51	-12	33	2595	71	4119	235	300	69	4057	93	630	77	507	89	191	90	4022	120	4331	284
d\WFData_Tube	10	10	10/6/15	124	186	28	1769	53	15	33	2321	70	3714	254	297	69	4136	93	557	78	596	90	272	90	3785	119	5068	315
d\WFData_Tube	10	10	10/6/15	124.5	115	28	1782	53	-14	32	2141	68	3914	241	277	70	3582	90	657	77	500	89	103	91	4150	121	4559	291
d\WFData_Tube	10	10	10/6/15	125	121	32	1766	53	-6	34	2479	71	4016	242	339	69	4024	92	571	77	545	90	160	91	4139	121	4596	294
d\WFData_Tube	10	10	10/6/15	125.5	123	29	1625	53	-17	35	2549	72	3686	240	338	69	4577	96	522	77	714	91	234	91	4025	121	4943	294
d\WFData_Tube	10	10	10/6/15	126	74	29	1776	55	-30	35	2289	69	3920	247	304	70	3919	94	708	79	551	91	115	92	4350	123	4733	301
d\WFData_Tube	10	10	10/6/15	126.5	129	31	1769	55	-6	35	2597	72	4008	239	190	70	4137	95	600	78	592	89	107	89	4322	121	4751	290
d\WFData_Tube	10	10	10/6/15	127	118	30	1987	57	-17	35	2495	72	3930	250	286	72	3966	94	750	78	603	91	134	91	4417	124	4905	310
d\WFData_Tube	10	10	10/6/15	127.5	178	30	2040	57	-21	34	1966	66	3906	239	438	70	3178	88	724	79	572	90	130	90	4676	124	5009	287
d\WFData_Tube	10	10	10/6/15	128	140	30	2041	57	-1	35	2289	69	3472	245	299	71	3951	93	678	78	812	92	129	91	4835	125	4430	302
d\WFData_Tube	10	10	10/6/15	128.5	156	30	2159	59	-46	36	2136	70	4038	253	509	73	3862	93	692	78	582	90	189	92	4612	125	5270	313
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d\WFData_Tube	10	10	10/6/15	129.5	214	30	2264	59	-14	33	1786	65	4545	253	388	72	2946	88	730	79	490	90	76	92	4896	127	4415	305
d\WFData_Tube	10	10	10/6/15	130	180	31	2386	60	37	33	1153	59	4397	264	567	72	2103	80	697	80	607	91	260	93	5040	129	4672	314
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d\WFData_Tube	10	10	10/6/15	131	210	30	2313	59	20	34	2138	68	4104	248	197	73	3764	94	738	79	636	91	203	92	4864	126	4765	304
d\WFData_Tube	10	10	10/6/15	131.5	175	31	2229	58	9	34	2613	72	4267	239	246	71	4682	99	602	79	524	90	283	90	4577	124	4578	294
d\WFData_Tube	10	10	10/6/15	132	203	32	2572	62	5	33	1595	63	4395	259	593	73	2502	84	622	79	543	90	241	91	5408	129	4755	314
d\WFData_Tube	10	10	10/6/15	132.5	183	32	2577	61	33	33	1686	62	4193	262	519	73	2442	84	805	80	533	91	180	93	5322	130	4886	324
d\WFData_Tube	10	10	10/6/15	133	234	31	2658	62	23	32	1457	61	4254	263	435	73	2220	83	708	80	681	93	149	93	5486	131	5214	321
d\WFData_Tube	10	10	10/6/15	133.5																								

45MM Data_7	10	10	10/7/15	41	148	31	2536	45	2.41	10	31	482	51	3711	279	474	72	690	46	10.01	981	78	667	90	215	92	5541	130	5645	138
45MM Data_7	10	10	10/7/15	41.5	150	32	2457	40	2.44	12	31	627	51	4204	277	845	72	618	46	10.68	758	78	471	90	263	92	5751	131	4956	138
45MM Data_7	10	10	10/7/15	42	155	31	2263	39	2.40	14	32	492	50	4396	280	850	72	684	46	9.85	788	79	618	91	177	92	5394	130	4615	138
45MM Data_7	10	10	10/7/15	42.5	173	32	2455	45	2.47	1	32	525	51	4068	283	825	73	514	46	12.84	785	78	679	91	148	92	5254	130	4948	138
45MM Data_7	10	10	10/7/15	43	140	33	2413	41	2.53	9	33	409	51	3738	289	894	74	712	47	9.41	949	80	631	92	214	93	5784	132	5480	132
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45MM Data_7	10	10	10/7/15	44	134	32	2380	40	2.52	0	31	471	51	4261	295	828	74	694	46	9.80	889	80	599	92	177	93	5430	132	5039	134
45MM Data_7	10	10	10/7/15	44.5	167	31	2395	40	2.51	13	31	452	50	3818	286	857	75	529	47	12.67	975	81	621	94	264	95	6053	133	5399	144
45MM Data_7	10	10	10/7/15	45	205	33	2264	39	2.40	13	32	470	50	4043	296	902	75	589	47	11.38	805	80	582	91	177	92	5887	131	5321	149
45MM Data_7	10	10	10/7/15	45.5	211	31	2402	39	2.46	1	31	489	50	4378	272	764	73	641	47	10.45	803	79	475	91	109	93	5886	132	4864	132
45MM Data_7	10	10	10/7/15	46	134	30	2433	40	2.46	10	31	614	51	4247	272	679	72	511	47	12.36	898	79	620	91	189	93	5652	131	4739	134
45MM Data_7	10	10	10/7/15	46.5	140	30	2646	42	2.34	31	30	565	50	4576	285	655	73	609	46	10.84	789	79	628	91	144	92	5394	129	5035	141
45MM Data_7	10	10	10/7/15	47	228	30	2502	40	2.40	-1	31	477	51	4073	270	580	72	523	46	12.82	714	78	752	91	132	92	5334	130	5286	123
45MM Data_7	10	10	10/7/15	47.5	148	30	2544	40	2.36	19	30	597	50	4532	267	572	72	687	48	9.30	915	78	626	92	143	94	5225	130	5024	139
45MM Data_7	10	10	10/7/15	48	131	31	2541	42	2.42	-21	32	537	51	4209	277	591	73	541	47	12.38	782	79	544	90	178	92	5132	129	5015	133
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45MM Data_7	10	10	10/7/15	49	142	30	2508	45	2.43	0	32	564	51	4268	278	641	72	636	48	10.69	791	79	657	91	136	93	5067	129	5188	129
45MM Data_7	10	10	10/7/15	49.5	170	32	2639	42	2.35	-10	31	540	51	4511	287	700	72	626	47	10.70	770	79	734	92	260	93	5290	130	4887	143
45MM Data_7	10	10	10/7/15	50	196	32	2502	45	2.44	-13	33	609	52	4294	292	726	72	556	47	12.05	771	80	623	92	133	92	5143	130	5334	149
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45MM Data_7	10	10	10/7/15	51	145	31	2623	42	2.36	5	32	612	54	4297	279	676	73	1017	72	7.08	838	79	688	92	107	92	4952	129	5000	137
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45MM Data_7	10	10	10/7/15	54	181	32	2466	40	2.43	44	31	562	51	4814	287	847	73	626	47	10.70	815	79	644	92	78	93	5559	131	4943	137
45MM Data_7	10	10	10/7/15	54.5	148	29	2597	45	2.35	-5	31	532	50	4035	274	740	72	769	48	8.84	634	78	642	91	20	92	5269	129	5184	139
45MM Data_7	10	10	10/7/15	55	176	31	2617	45	2.48	-7	33	420	50	4381	277	607	71	717	47	9.29	813	79	553	90	168	93	5331	130	5099	125
45MM Data_7	10	10	10/7/15	55.5	177	32	2538	42	2.44	-4	33	642	52	3995	292	473	72	674	48	10.89	840	79	706	93	144	94	6029	134	5389	154
45MM Data_7	10	10	10/7/15	56	183	30	2511	39	2.31	47	29	603	51	4399	267	693	71	617	47	10.86	643	78	752	91	151	93	6095	133	4712	159
45MM Data_7	10	10	10/7/15	56.5	133	31	2542	45	2.39	29	30	655	51	4275	291	721	72	671	47	9.99	627	80	728	93	194	94	6188	134	5254	159
45MM Data_7	10	10	10/7/15	57	211	31	2497	45	2.44	-12	31	754	53	3817	270	600	71	636	46	10.41	697	78	554	91	155	93	6861	137	5367	129
45MM Data_7	10	10	10/7/15	57.5	179	30	2570	45	2.37	27	31	675	52	4370	274	548	73	548	47	12.23	883	78	544	92	69	93	6844	136	5144	132
45MM Data_7	10	10	10/7/15	58	183	31	2791	43	2.26	19	31	666	53	4246	278	639	73	553	47	12.22	818	78	651	91	170	92	4286	134	5132	132
45MM Data_7	10	10	10/7/15	58.5	145	32	2713	42	2.29	25	31	609	51	4489	276	674	72	686	48	9.91	740	79	796	93	178	92	4286	134	4976	132
45MM Data_7	10	10	10/7/15	59	222	32	2581	42	2.40	-4	32	794	55	4204	275	795	72	1242	71	6.22	942	80	688	92	214	94	5823	132	5365	131
45MM Data_7	10	10	10/7/15	59.5	176	31	2642	42	2.33	-13	32	598	52	4530	268	676	73	851	69	8.11	852	80	497	91	113	93	5752	131	4759	138
45MM Data_7	10	10	10/7/15	60	179	31	2750	43	2.29	-4	31	594	53	4317	277	593	73	625	70	11.20	872	81	680	92	175	94	5877	134	4864	138
45MM Data_7	10	10	10/7/15	60.5	216	32	2694	42	2.30	-4	31	558	52	4354	285	765	73	654	48	10.40	834	81	587	92	119	93	5457	132	5205	148
45MM Data_7	10	10	10/7/15	61	216	32	2711	43	2.31	28	32	549	50	4280	285	595	72	673	47	9.36	817	79	755	91	180	90	5445	131	4725	141
45MM Data_7	10	10	10/7/15	61.5	184	32	2776	44	2.31	25	33	642	52	4648	280	643	73	566	48	12.01	812	80	641	92	199	92	5657	131	4853	138
45MM Data_7	10	10	10/7/15	62	156	32	3044	43	2.34	20	32	600	52	4742	278	694	72	640	47	10.47	806	80	796	92	64	94	5863	133	4947	138
45MM Data_7	10	10	10/7/15	62.5	147	31	2838	43	2.24	-2	31	625	52	4590	276	724	74	718	48	9.47	763	80	448	91	94	93	6028	134	5289	137
45MM Data_7	10	10	10/7/15	63	198	31	2882	43	2.34	-2	33	492	51	4712	280	584	73	666	49	10.36	811	81	655	92	300	94	5846	132	4912	133
45MM Data_7	10	10	10/7/15	63.5	184	30	2639	42	2.35	-8	33	535	53	4087	274	648	72	824	48	8.25	743	78	509	91	36	92	5425	131	5484	137
45MM Data_7	10	10	10/7/15	64	217	31	2859	44	2.24	23	32	585	52	4587	266	708	72	606	47	11.06	920	80	639	92	189	92	5989	132	4959	138
45MM Data_7	10	10	10/7/15	64.5	163	33	2758	44	2.32	0	34	535	53	4374	292	740	74	704	49	9.80	800	80	578	92	146	93	5753	132	4774	131
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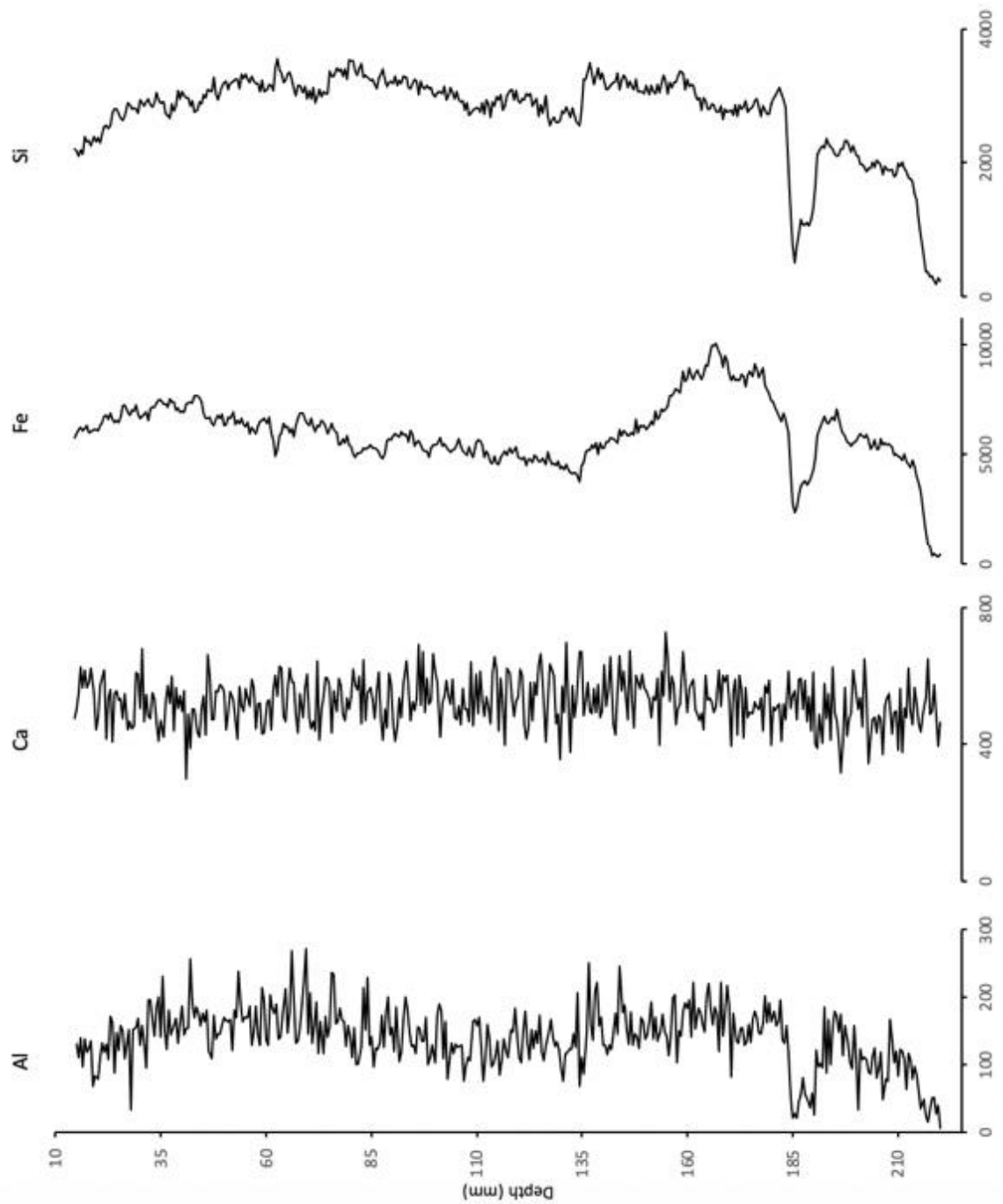
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45MM Data_1	10	10	10/7/15	86.5	245	33	3311	46	2.34	-9	32	598	52	4407	277	794	76	732	69	9.43	919	81	570	93	54	94	4528	136	5211	343
45MM Data_1	10	10	10/7/15	87	217	34	3269	46	2.11	-12	35	477	53	4388	278	748	76	633	69	10.90	1000	82	626	93	112	94	4348	137	5071	343
45MM Data_1	10	10	10/7/15	87.5	178	33	3248	46	2.09	-21	33	640	52	4414	272	811	73	825	68	8.24	817	80	700	93	117	94	4631	137	5263	334
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45MM Data_1	10	10	10/7/15	90	280	32	3071	45	2.12	-4	31	582	51	4398	264	879	75	611	67	10.97	753	79	508	92	80	94	4783	141	5485	321
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45MM Data_1	10	10	10/7/15	91	213	33	3015	46	2.19	-1	33	558	52	4604	277	804	74	668	68	10.38	909	81	593	92	89	93	4952	137	5042	335
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45MM Data_1	10	10	10/7/15	94	168	33	3240	47	2.13	-2	33	538	53	4876	269	884	75	614	67	10.91	811	80	658	92	213	93	4840	139	4951	324
45MM Data_1	10	10	10/7/15	94.5	205	34	3390	49	2.06	-1	33	494	52	4972	276	897	74	653	67	10.26	914	81	571	90	132	90	4814	137	4916	333
45MM Data_1	10	10	10/7/15	95	179	32	3389	48	2.01	4	32	482	52	4839	267	895	75	497	66	13.28	969	80	754	92	142	94	4596	137	5554	331
45MM Data_1	10	10	10/7/15	95.5	234	34	3432	49	2.01	0	32	582	52	4311	270	1006	74	745	69	9.26	923	82	686	92	273	94	4842	138	5351	333
45MM Data_1	10	10	10/7/15	96	176	33	3404	70	2.06	-9	35	492	52	4797	264	818	75	544	68	12.50	928	81	628	94	80	94	4887	139	5282	343
45MM Data_1	10	10	10/7/15	96.5	207	35	3245	49	2.13	-9	34	498	51	4894	277	842	75	511	68	12.34	1046	82	628	92	86	94	4788	137	5408	342
45MM Data_1	10	10	10/7/15	97	194	34	3447	70	2.03	22	33	586	52	4554	277	892	76	819	68	8.80	833	80	537	92	112	94	4660	137	5173	341
45MM Data_1	10	10	10/7/15	97.5	238	33	3537	70	1.99	-27	32	489	52	4538	277	841	75	637	68	10.68	978	81	619	92	176	94	4630	138	5274	342
45MM Data_1	10	10	10/7/15	98	258	34	3456	69	2.00	-9	32	550	52	5013	275	851	75	638	68	10.66	848	81	609	91	296	93	4719	140	4129	332
45MM Data_1	10	10	10/7/15	98.5	203	34	3342	69	2.06	-22	34	487	52	4608	282	771	76	612	68	11.31	889	81	509	92	94	93	4854	139	5334	343
45MM Data_1	10	10	10/7/15	99	247	33	3545	70	1.97	-15	33	500	52	4435	274	870	76	564	68	12.06	834	81	569	91	110	92	4721	138	5275	339
45MM Data_1	10	10	10/7/15	99.5	201	33	3549	70	1.97	-1	33	572	52	4319	270	841	76	439	68	15.49	881	81	528	91	139	94	4977	139	5185	336
45MM Data_1	10	10	10/7/15	200	207	33	3446	70	2.03	-11	34	565	53	4516	270	984	75	733	68	9.41	879	80	531	91	216	93	4763	139	5756	334
45MM Data_1	10	10	10/7/15	100.5	242	34	3378	69	2.04	-23	33	542	53	4260	280	992	76	702	68	9.69	901	81	656	91	71	93	4788	140	5499	340
45MM Data_1	10	10	10/7/15	201	239	33	3373	69	2.05	24	33	529	52	4684	286	1034	76	621	69	11.21	929	81	647	92	293	94	4769	140	5398	352
45MM Data_1	10	10	10/7/15	101.5	214	33	3492	70	2.00	-4	34	541	53	4352	284	834	75	656	68	10.37	1172	82	738	92	117	93	4787	140	5190	352
45MM Data_1	10	10	10/7/15	202	243	35	3429	69	2.01	2	32	589	53	4611	275	934	75	695	68	9.78	854	81	623	93	146	93	4950	140	5342	339
45MM Data_1	10	10	10/7/15	202.5	297	34	3379	68	2.01	37	31	679	54	4895	274	809	76	693	70	10.10	1094	81	611	93	34	94	4715	140	5250	336
45MM Data_1	10	10	10/7/15	203	233	34	3534	70	1.98	18	33	586	52	4546	269	859	76	726	68	8.37	901	80	555	92	131	94	4765	141	5321	336
45MM Data_1	10	10	10/7/15	203.5	184	34	3410	69	2.02	-5	33	546	53	4268	281	867	76	607	69	11.37	1085	81	681	92	253	95	4671	140	4799	349
45MM Data_1	10	10	10/7/15	204	167	33	3408	70	2.04	10	34	529	53	4400	279	866	76	662	69	10.42	915	81	567	93	88	95	4847	142	5404	345
45MM Data_1	10	10	10/7/15	204.5	289	34	3466	70	2.02	-19	35	446	52	5133	270	875	77	552	70	12.68	986	81	597	92	280	94	4710	140	4984	329
45MM Data_1	10	10	10/7/15	205	244	35	3452	70	2.03	15	35	637	53	4517	276	979	76	653	68	10.41	833	81	601	92	254	94	4950	139	5385	338
45MM Data_1	10	10	10/7/15	205.5	177	33	3384	71	1.98	14	34	607	52	4352	273	947	74	788	69	8.76	908	80	486	92	260	93	4727	139	5403	338
45MM Data_1	10	10	10/7/15	206	246	34	3468	70	2.02	-4	33	628	54	4638	283	956														

45000Data_7	30	30	10/7/15	123	399	33	2900	66	2.28	-0.6	35	1314	41	4821	274	458	75	2838	86	8.03	938	85	426	92	93	93	4463	134	4797	334
45000Data_7	30	30	10/7/15	123.5	399	34	3402	70	2.06	-0.9	33	629	53	4603	286	705	76	838	71	8.57	919	82	513	92	340	94	7158	340	5497	353
45000Data_7	30	30	10/7/15	124	387	34	3428	70	2.04	1	34	192	53	3507	305	814	77	598	70	11.71	941	82	733	93	309	94	7547	341	6405	383
45000Data_7	30	30	10/7/15	124.5	388	33	3390	68	2.06	30	34	402	52	4229	273	780	76	715	69	9.45	835	85	763	92	235	92	6893	338	5547	334
45000Data_7	30	30	10/7/15	125	377	33	3217	67	2.08	20	32	636	53	4378	284	830	76	584	68	11.64	1018	85	478	92	252	94	6926	338	5807	348
45000Data_7	30	30	10/7/15	125.5	378	34	3394	69	2.03	9	34	518	52	4283	279	819	76	615	69	11.22	1009	85	652	92	365	94	6979	339	5178	345
45000Data_7	30	30	10/7/15	126	341	35	3383	70	2.07	-26	35	942	54	4263	287	884	76	687	70	10.19	919	83	632	94	214	93	7338	340	5890	355
45000Data_7	30	30	10/7/15	126.5	368	33	3388	68	2.05	0	34	586	53	4867	306	943	77	638	70	10.87	872	80	534	93	358	95	7121	340	4915	323
45000Data_7	30	30	10/7/15	127	399	34	3398	69	2.03	24	33	647	53	4490	271	1072	76	745	69	9.26	856	80	602	93	332	95	7207	339	5698	333
45000Data_7	30	30	10/7/15	127.5	347	35	3423	70	2.05	9	33	768	55	4732	292	975	76	838	70	8.35	932	82	767	94	385	94	7161	339	5698	361
45000Data_7	30	30	10/7/15	128	292	33	3466	69	1.99	26	33	679	53	4997	289	865	76	627	69	11.80	908	82	613	93	317	93	6874	339	4799	327
45000Data_7	30	30	10/7/15	128.5	245	33	3326	69	2.07	6	35	642	54	4882	288	907	76	707	70	9.80	901	80	595	92	381	93	6827	337	4929	325
45000Data_7	30	30	10/7/15	129	380	33	3220	68	2.11	-17	35	760	55	4580	282	870	75	999	72	7.21	1085	85	512	92	83	94	7178	339	5034	309
45000Data_7	30	30	10/7/15	129.5	388	33	3504	70	2.00	20	33	705	54	4911	283	958	75	747	69	9.24	715	80	618	92	314	93	6620	337	5395	323
45000Data_7	30	30	10/7/15	130	300	33	3404	70	2.06	16	34	606	52	4731	270	895	75	585	69	11.79	814	80	637	92	240	92	7202	339	5196	329
45000Data_7	30	30	10/7/15	130.5	232	33	3270	68	2.08	62	33	666	52	4830	288	935	76	681	69	10.13	1008	82	610	93	205	94	7047	340	5333	329
45000Data_7	30	30	10/7/15	131	294	34	3271	67	2.15	-3	33	507	52	4491	288	1079	78	709	68	9.47	1072	85	601	93	306	93	7396	340	5103	344
45000Data_7	30	30	10/7/15	131.5	239	35	3239	69	2.13	-0.9	34	508	53	4886	298	1007	77	620	69	11.53	940	82	502	91	302	92	7145	338	5465	325
45000Data_7	30	30	10/7/15	132	219	33	2905	65	2.19	4	32	632	52	4428	271	957	76	675	69	10.23	796	80	543	91	58	93	6898	339	5363	325
45000Data_7	30	30	10/7/15	132.5	246	32	3090	65	2.17	-9	32	637	53	4967	279	821	76	633	69	10.80	1138	82	652	91	98	94	7088	339	4908	336
45000Data_7	30	30	10/7/15	133	245	33	2905	65	2.24	9	33	549	52	4377	281	895	77	702	70	9.87	905	82	645	94	417	95	6968	338	5289	349
45000Data_7	30	30	10/7/15	133.5	237	32	2970	65	2.19	9	31	611	52	4283	280	787	76	718	69	9.41	941	85	578	93	245	94	7086	338	5030	343
45000Data_7	30	30	10/7/15	134	272	32	3036	65	2.14	0	32	540	52	4290	277	1015	75	761	67	9.56	944	85	617	93	144	94	6812	339	5437	343
45000Data_7	30	30	10/7/15	134.5	230	35	2979	66	2.22	-62	33	935	52	4511	288	1035	75	592	67	11.32	863	85	684	93	227	93	6889	339	5205	349
45000Data_7	30	30	10/7/15	135	236	34	2836	65	2.29	6	33	504	52	4475	271	787	74	604	67	11.09	919	80	482	91	76	93	6821	338	5136	323
45000Data_7	30	30	10/7/15	135.5	196	34	2970	66	2.22	8	33	611	53	4640	285	963	77	596	68	11.41	937	85	649	92	206	93	6966	338	5246	309
45000Data_7	30	30	10/7/15	136	181	33	2969	66	2.21	15	33	407	51	4462	278	995	76	572	67	11.71	857	80	525	91	348	92	7105	338	5357	334
45000Data_7	30	30	10/7/15	136.5	230	32	3003	65	2.18	-11	32	505	51	4652	276	830	76	544	69	12.68	1037	85	682	93	308	93	7188	338	5141	344
45000Data_7	30	30	10/7/15	137	199	32	3099	66	2.13	15	32	624	52	4442	270	1017	76	585	68	11.62	767	80	629	92	136	93	6938	338	5188	329
45000Data_7	30	30	10/7/15	137.5	287	35	3024	66	2.18	3	33	438	52	4217	284	1117	75	593	67	11.80	936	85	515	91	96	92	7159	337	5813	343
45000Data_7	30	30	10/7/15	138	246	33	3126	67	2.14	-17	32	630	53	4412	284	868	76	663	69	10.41	899	85	536	92	319	94	6817	338	5197	363
45000Data_7	30	30	10/7/15	138.5	245	34	2968	65	2.20	25	31	674	53	4479	285	767	75	629	69	10.87	900	80	641	92	344	92	6871	337	5353	347
45000Data_7	30	30	10/7/15	139	205	34	3022	66	2.18	-5	33	620	53	4294	287	824	76	679	69	10.16	1057	82	619	93	91	94	6563	336	5001	353
45000Data_7	30	30	10/7/15	139.5	238	34	2966	65	2.19	4	32	624	53	4875	284	738	75	789	70	8.87	817	80	408	92	240	93	6881	336	5409	344
45000Data_7	30	30	10/7/15	140	263	32	2740	62	2.26	33	31	881	55	4650	280	734	76	1077	74	6.87	830	85	652	92	332	93	6438	336	5365	347
45000Data_7	30	30	10/7/15	140.5	177	32	2787	65	2.13	-17	34	709	55	4507	290	846	75	892	71	7.96	830	80	630	93	294	93	6808	336	5345	340
45000Data_7	30	30	10/7/15	141	200	33	2862	64	2.24	12	31	586	52	4375	287	820	75	742	68	9.16	782	80	731	93	386	92	6265	335	5229	349
45000Data_7	30	30	10/7/15	141.5	263	32	2983	65	2.18	2	31	717	54	4327	286	847	75	725	69	9.52	932	80	613	92	314	94	6246	336	5399	349
45000Data_7	30	30	10/7/15	142	223	32	3078	66	2.14	1	33	560	53	4225	280	756	74	682	68	9.97	869	80	489	91	245	93	6280	335	5430	343
45000Data_7	30	30	10/7/15	142.5	214	33	2925	65	2.22	-44	34	682	54	4826	274	807	74	615	68	11.36	1040	85	778	92	71	92	6399	335	4616	345
45000Data_7	30	30	10/7/15	143	219	34	2935	64	2.26	4	32	666	53	4414	278	869	75	654	69	10.55	839	80	534	92	335	94	6596	337	4953	338
45000Data_7	30	30	10/7/15	143.5	262	34	2857	65	2.28	-25	33	618	54	4404	281	1010	74	869	69	7.94	868	81	695	93	140	93	6452	335	5213	339
45000Data_7	30	30	10/7/15	144	232	32	2934	65	2.22	10	32	568	52	4538	269	785	75	823	68	8.26	921	81	601	93	250	95	6344	335	4996	336
45000Data_7	30	30	10/7/15	144.5	197	32	2806	64	2.28	4	33	539	53	4389	270	714	74	732	67	9.15	757	79	674	92	288	92	6286	335	5217	342
45000Data_7	30	30	10/7/15	145	239	31	2870	63	2.30	10	31	629	53	4371	280	581	74	538	68	13.13	830	80	630	92	288	92	6475	335	5228	342
45000Data_7	30	30	10/7/15	145.5	274	33	2913	64	2.30	4	31	614	53	4914	274	811	75	573	68	11.87	802	79	824	92	183	91	6374	334	4941	329
45000Data_7	30	30	10/7/15	146	210	33	2749	63	2.29	0	32	766	54	4587	274	840	74	882	70	7.94	852	79	595	92	203	93	6234	334	4964	329
45000Data_7	30	30	10/7/15	146.5	173	31	2723	63	2.31	5	34	1018	57	4484	267	625	74	1452	71	5.31	960	79	554	92	197	93	5920	332	4805	

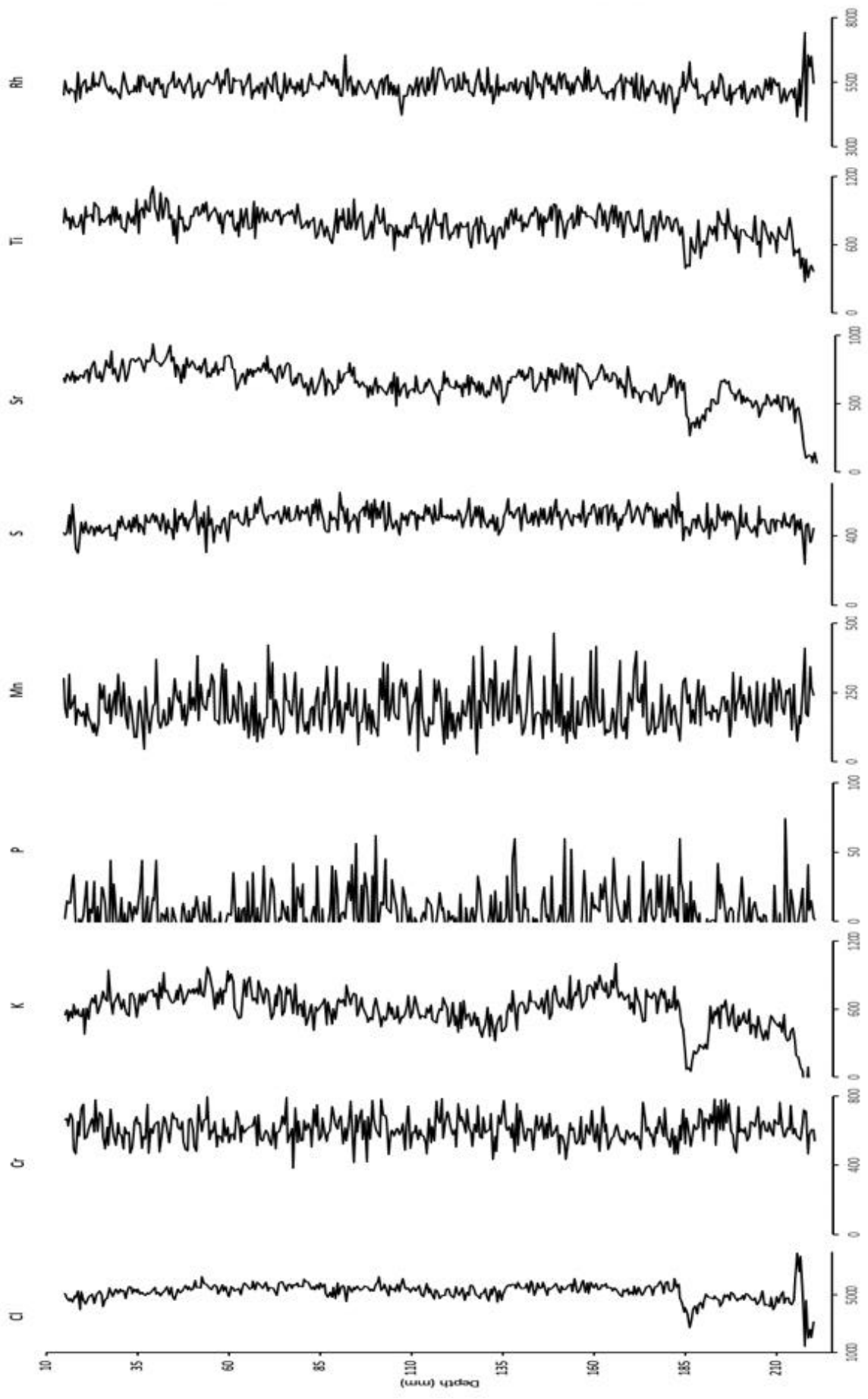
45MM Data_7	10	10	10/7/25	244	195	30	2211	57	2.85	16	31	636	53	34517	299	427	70	751	66	8.79	734	79	498	91	194	90	4930	127	5688	360
45MM Data_7	10	10	10/7/25	164.5	119	29	2042	55	2.87	17	29	636	51	4214	285	450	70	587	66	11.24	664	79	407	90	310	91	4793	125	4880	334
45MM Data_7	10	10	10/7/25	265	112	28	1897	54	2.85	17	31	615	52	3737	286	492	70	839	69	7.25	651	77	539	90	146	92	4535	124	5189	332
45MM Data_7	10	10	10/7/25	163.5	94	29	1977	55	2.78	17	29	638	52	4441	299	437	69	846	67	7.82	633	78	572	90	189	92	4562	125	4616	347
45MM Data_7	10	10	10/7/25	246	96	28	2912	55	2.76	-1	31	585	51	3934	295	561	68	734	65	8.86	757	79	589	90	179	92	4671	126	5025	343
45MM Data_7	10	10	10/7/25	166.5	128	28	1926	55	2.88	-7	32	570	52	4143	303	539	70	623	66	10.59	677	79	582	90	134	90	4861	126	4936	350
45MM Data_7	10	10	10/7/25	247	85	28	2017	56	2.75	-8	30	595	51	3879	294	333	70	689	67	9.72	655	78	454	90	309	91	4636	125	5073	347
45MM Data_7	10	10	10/7/25	167.5	85	29	1980	56	2.83	2	31	615	52	3980	309	308	70	740	67	9.05	740	79	544	90	17	91	4624	125	5137	346
45MM Data_7	10	10	10/7/25	248	122	28	1834	53	2.89	20	30	585	51	4213	290	447	69	796	66	8.29	616	78	619	91	146	92	4768	127	4754	333
45MM Data_7	10	10	10/7/25	168.5	147	28	1835	54	2.91	2	31	537	50	3733	295	476	69	635	66	10.39	635	78	549	91	237	93	4561	125	5120	335
45MM Data_7	10	10	10/7/25	249	122	29	1842	54	2.93	-5	30	568	51	4250	300	422	70	663	66	9.95	611	77	758	91	104	91	4570	125	4977	344
45MM Data_7	10	10	10/7/25	169.5	130	29	1809	54	2.99	0	32	640	52	4239	299	452	68	710	66	9.30	637	77	642	91	288	92	4433	124	4878	338
45MM Data_7	10	10	10/7/25	170	101	28	1888	54	2.86	17	29	592	50	4219	292	545	69	758	67	8.84	568	79	627	92	146	93	4551	125	4808	336
45MM Data_7	10	10	10/7/25	170.5	118	27	1847	53	2.87	13	30	588	51	3234	297	474	69	613	65	10.60	519	78	605	90	108	91	4452	125	5439	348
45MM Data_7	10	10	10/7/25	171	122	28	1860	54	2.90	1	31	609	53	4540	306	368	70	582	67	11.51	670	77	695	90	100	90	4678	125	4961	353
45MM Data_7	10	10	10/7/25	171.5	109	29	1738	53	3.07	-23	32	549	53	4890	298	380	69	851	68	7.99	621	77	631	91	210	92	5463	131	3888	335
45MM Data_7	10	10	10/7/25	172	69	28	1737	53	3.05	38	32	508	51	3382	321	395	69	780	66	8.46	661	79	611	92	154	93	4742	130	5793	375
45MM Data_7	10	10	10/7/25	172.5	144	27	1894	53	2.80	39	29	657	52	3958	284	399	70	753	67	8.90	648	78	591	91	299	91	4280	133	4833	330
45MM Data_7	10	10	10/7/25	173	124	29	1948	55	2.82	11	31	609	52	4337	287	459	71	618	67	10.84	655	78	749	92	133	91	5348	128	4678	331
45MM Data_7	10	10	10/7/25	173.5	87	29	1984	55	2.77	17	30	612	53	4264	288	485	70	704	68	8.90	581	78	648	91	242	92	4886	127	4602	338
45MM Data_7	10	10	10/7/25	174	117	28	1858	54	2.91	-4	32	608	53	3935	289	512	69	794	67	8.44	690	78	752	91	247	91	4213	132	4995	334
45MM Data_7	10	10	10/7/25	174.5	123	28	1820	55	2.86	14	31	580	52	3854	288	459	70	731	67	9.17	624	78	630	91	111	93	7081	137	5083	371
45MM Data_7	10	10	10/7/25	175	118	28	1814	54	2.98	-14	31	740	53	4238	290	473	69	807	66	8.18	658	77	641	91	164	90	4678	144	4617	334
45MM Data_7	10	10	10/7/25	175.5	143	29	1917	55	2.75	-7	30	722	52	3700	296	391	70	634	67	10.17	599	79	683	92	438	93	7603	139	5362	351
45MM Data_7	10	10	10/7/25	176	133	28	1863	54	2.90	47	32	672	52	3700	290	419	68	833	67	8.04	698	78	695	90	255	90	5398	129	5338	338
45MM Data_7	10	10	10/7/25	176.5	140	29	1907	54	2.83	27	31	570	52	4294	287	530	69	678	66	9.73	630	78	546	91	136	92	4943	128	4637	338
45MM Data_7	10	10	10/7/25	177	93	29	1931	55	2.85	0	32	522	52	3884	313	523	70	730	66	9.04	739	79	732	92	215	92	5545	129	5274	348
45MM Data_7	10	10	10/7/25	177.5	88	28	1790	55	2.96	-4	31	653	53	4394	304	369	70	552	66	11.96	780	78	577	92	118	92	5161	131	4669	348
45MM Data_7	10	10	10/7/25	178	101	27	1856	54	2.91	4	31	657	52	3963	279	406	69	697	67	9.61	629	79	603	90	141	93	5467	130	4719	322
45MM Data_7	10	10	10/7/25	178.5	147	30	1838	54	2.94	-80	32	726	54	3970	311	430	71	876	70	8.01	606	79	604	92	187	93	5149	128	5085	364
45MM Data_7	10	10	10/7/25	179	120	29	1834	55	3.00	-13	33	812	55	3908	295	500	70	1122	71	6.33	809	78	523	90	88	92	4997	128	5185	343
45MM Data_7	10	10	10/7/25	179.5	127	28	2214	55	2.80	31	28	919	54	4296	274	300	71	1022	71	6.95	795	80	681	92	184	93	5437	129	4127	328
45MM Data_7	10	10	10/7/25	180	119	29	2054	54	2.73	-4	31	840	55	4077	287	351	70	864	70	8.30	642	78	569	91	176	92	5489	129	4672	341

APPENDIX C

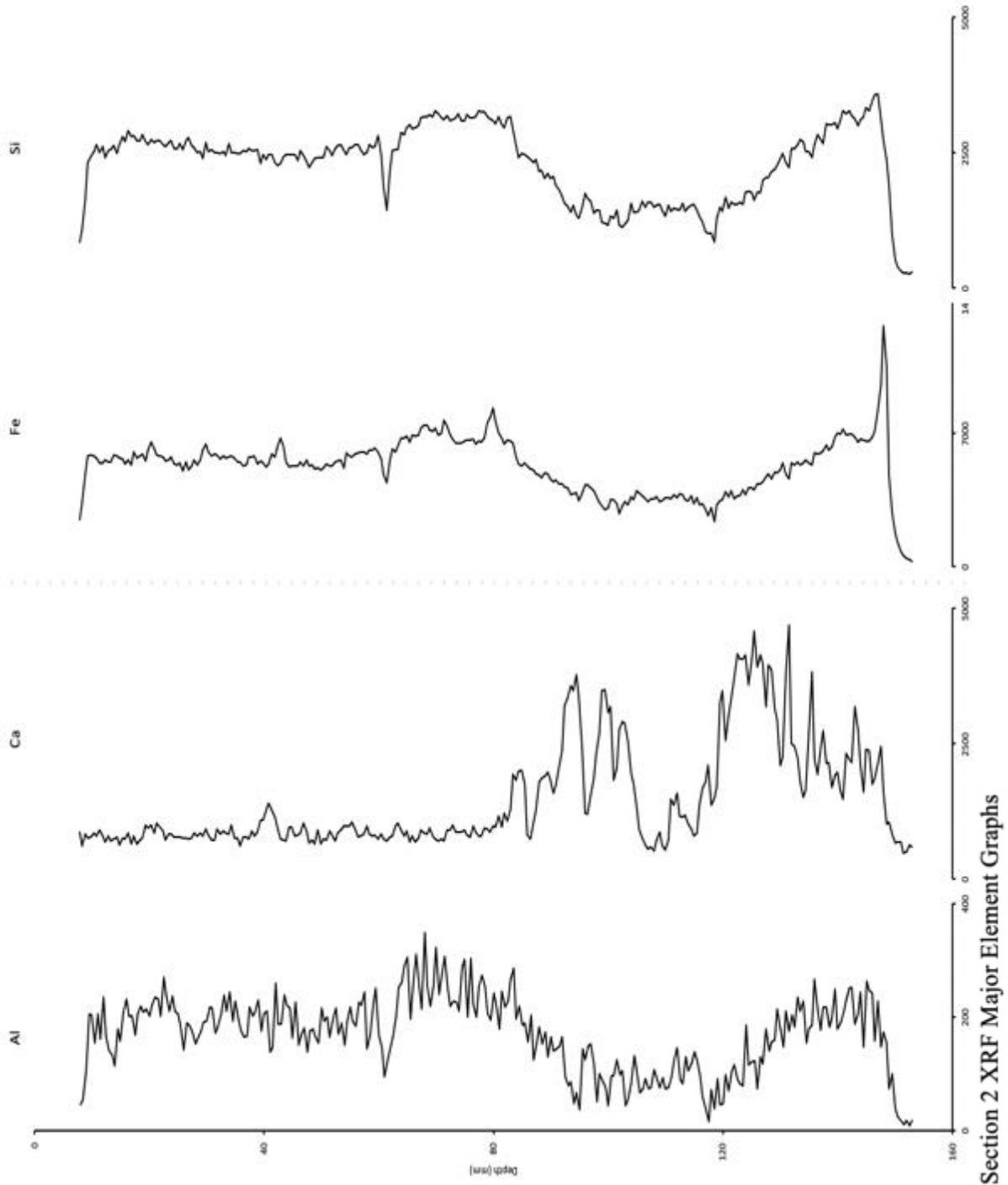
XRF Section Graphs of Major and Minor Elements



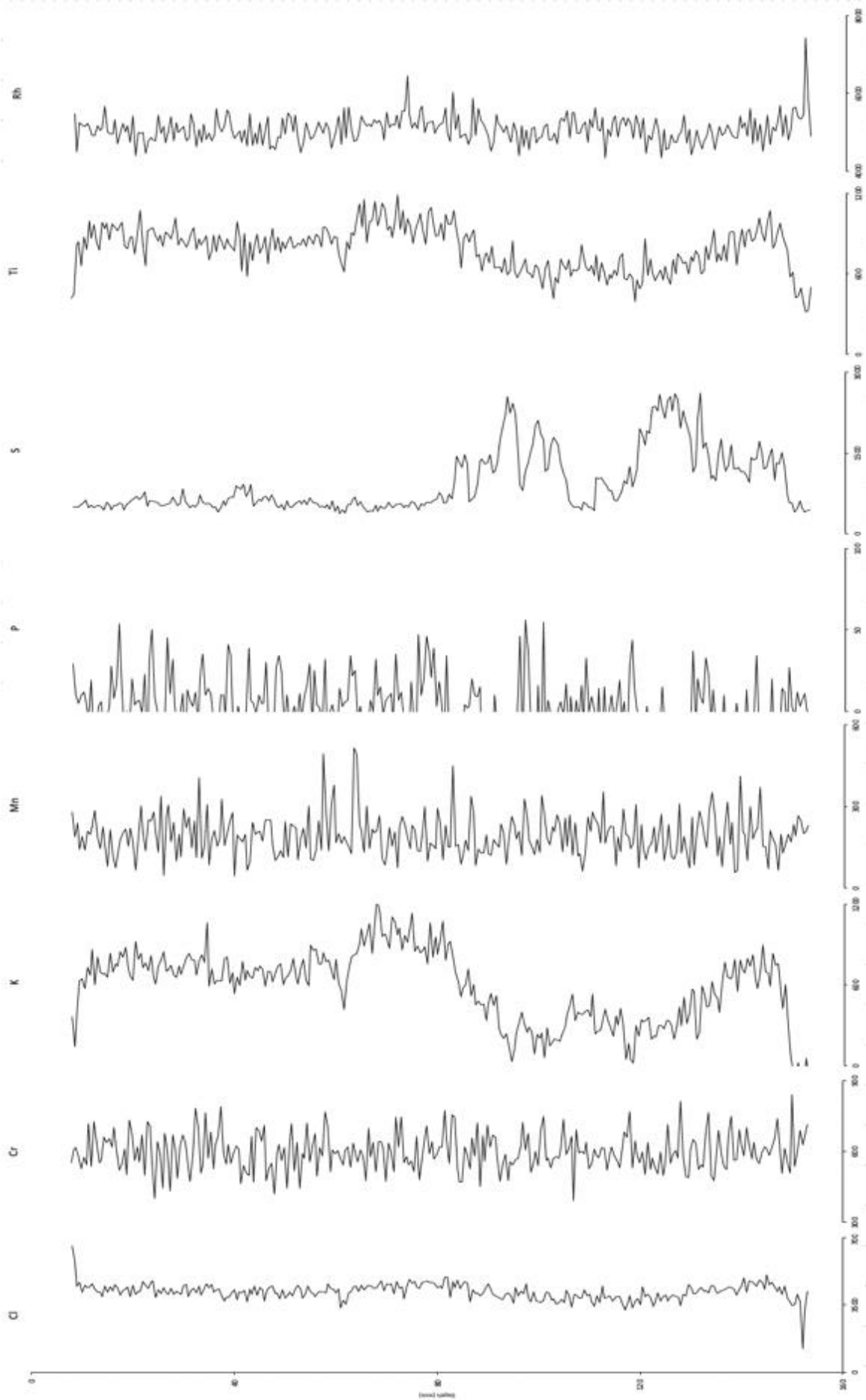
Section 1 XRF Major Element Graphs



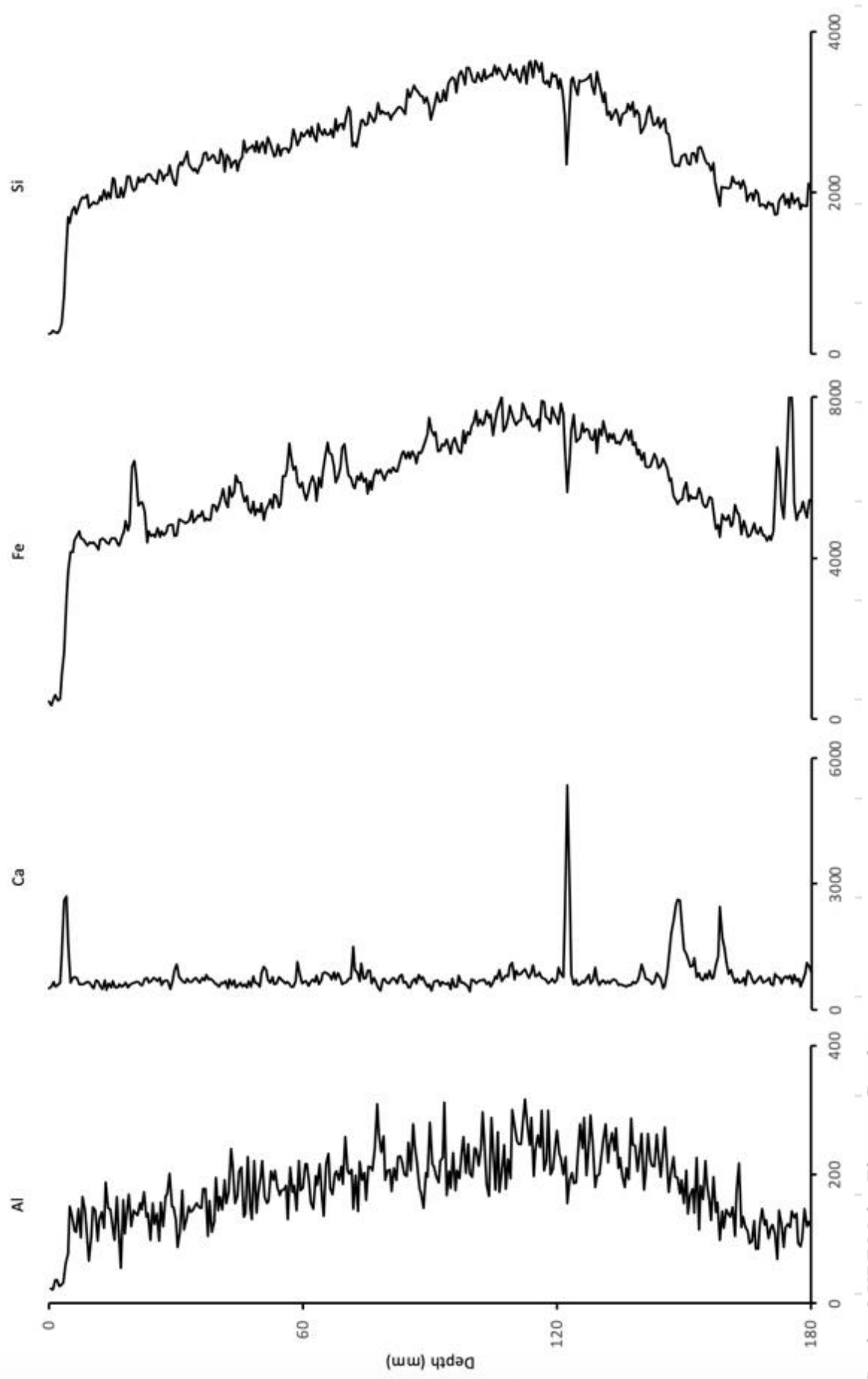
Section 1 XRF Minor Element Graphs



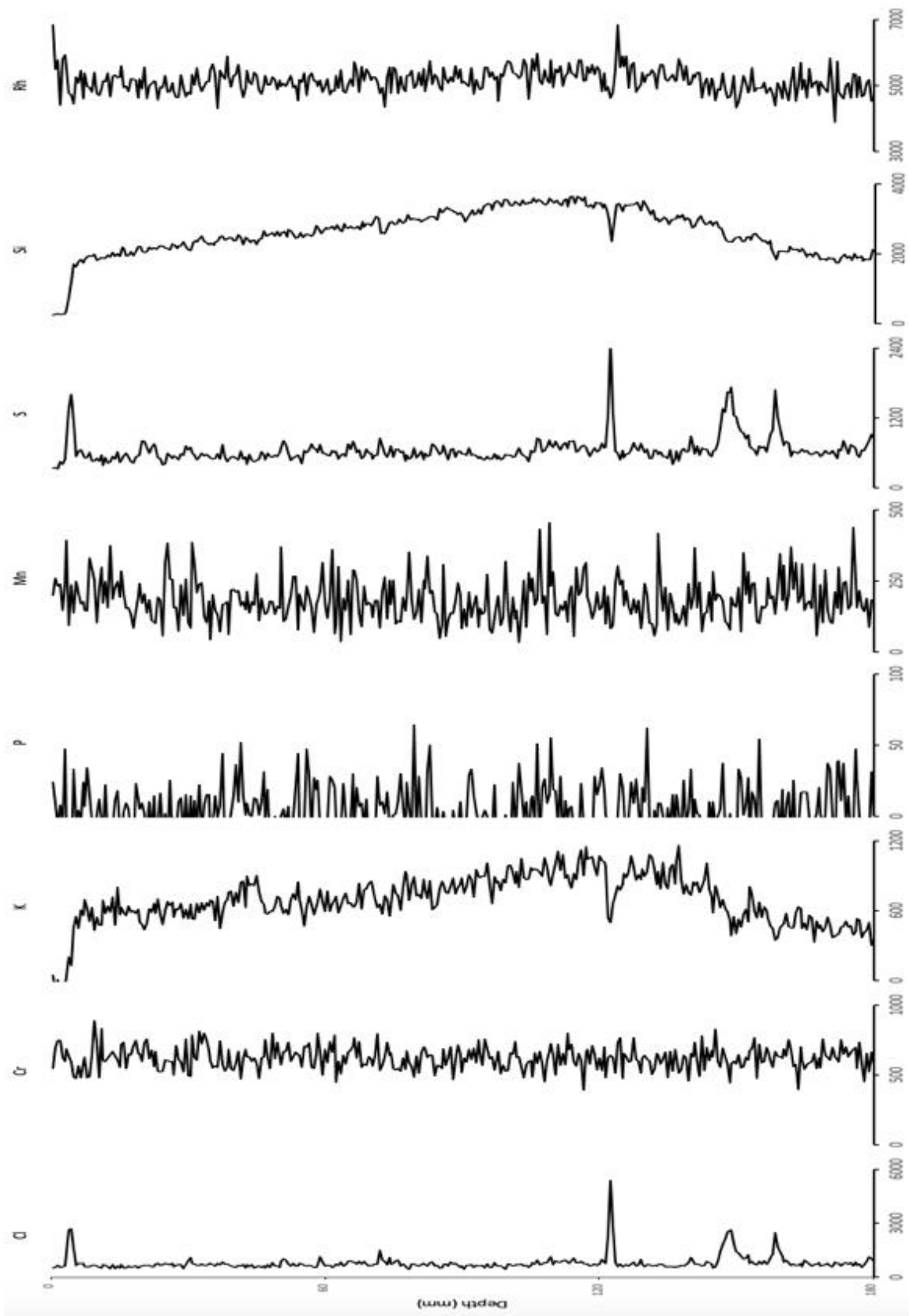
Section 2 XRF Major Element Graphs



Section 2 XRF Minor Element Graphs



Section 3 XRF Major Element Graphs

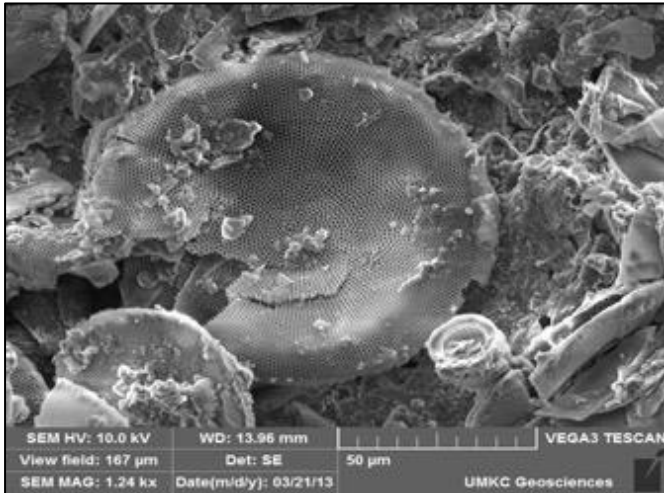


Section 3 XRF Minor Element Graphs

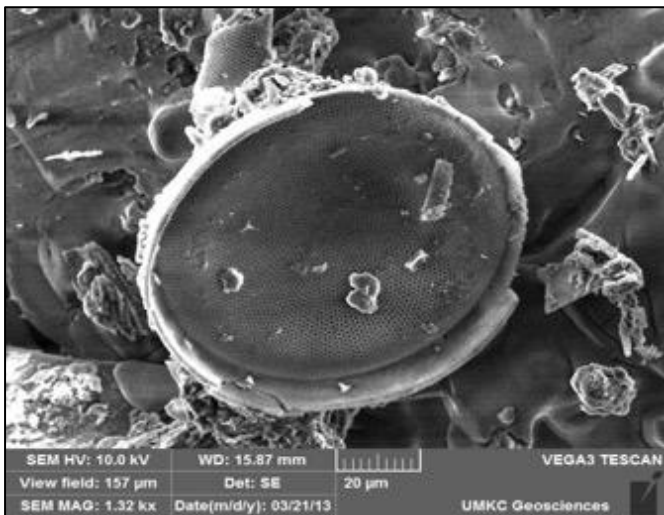
APPENDIX D

Scanning Electron Microscopy Diatom Images

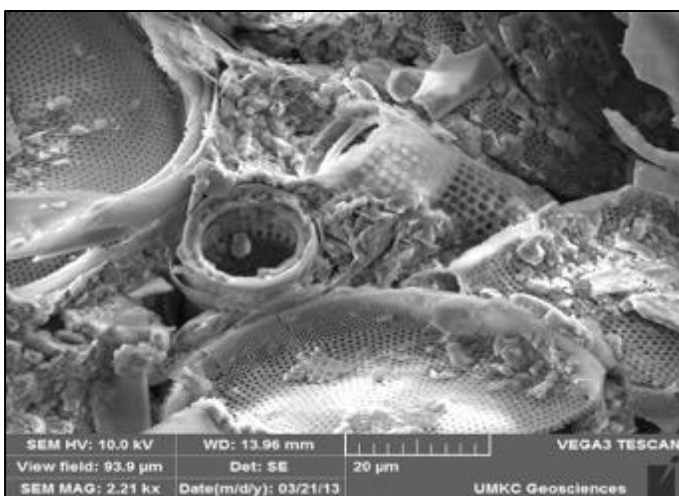
Section 1 (14 -14.50 m)



Stephanodiscus sp.

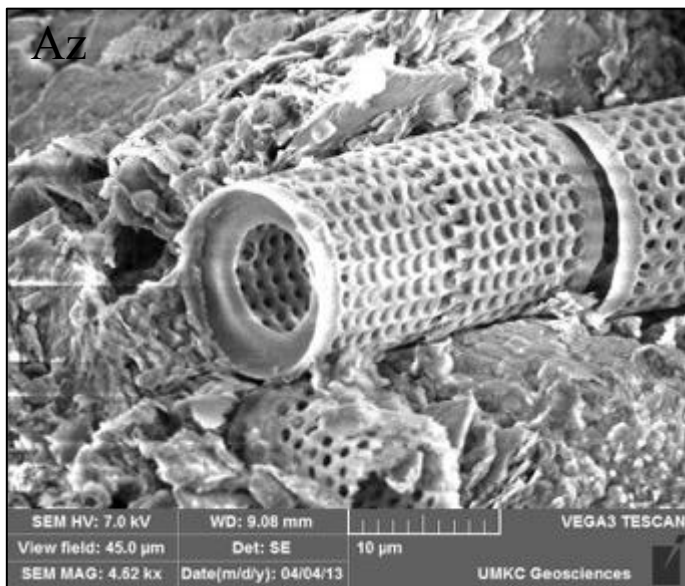


Stephanodiscus sp.

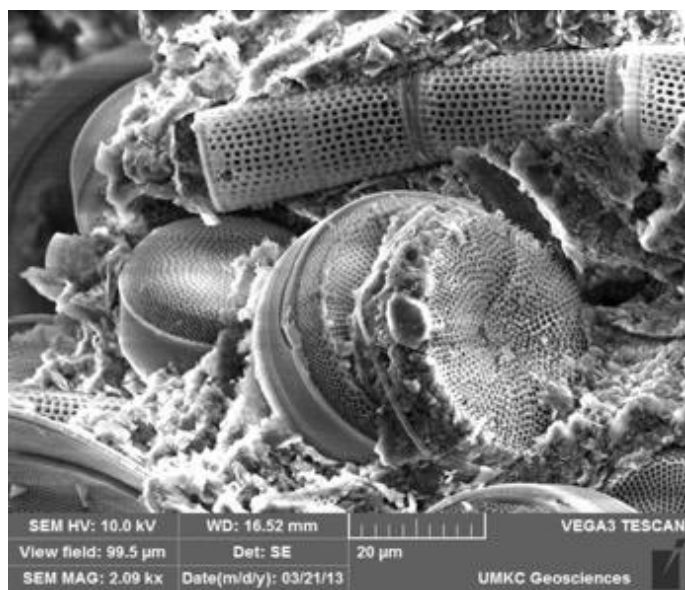


Stephanodiscus sp. and *Aulacoseira lirata*

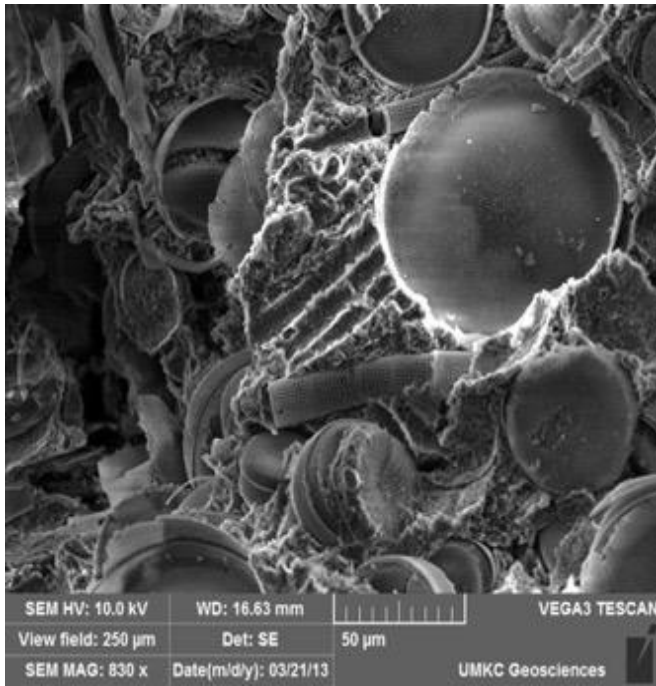
Section 2 (19.62 - 19.76 m)



Aulacoseira lirata

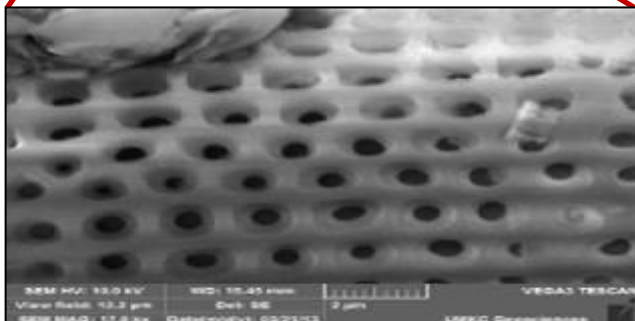
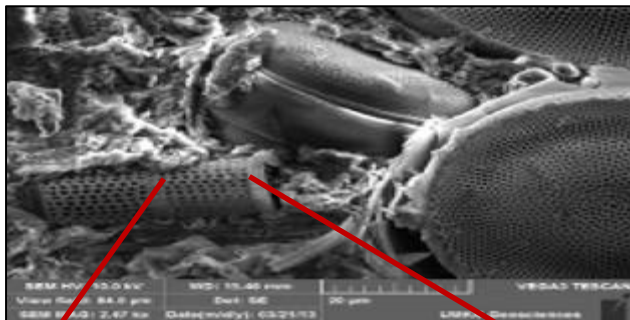


Aulacoseira lirata and *Stephanodiscus* sp.

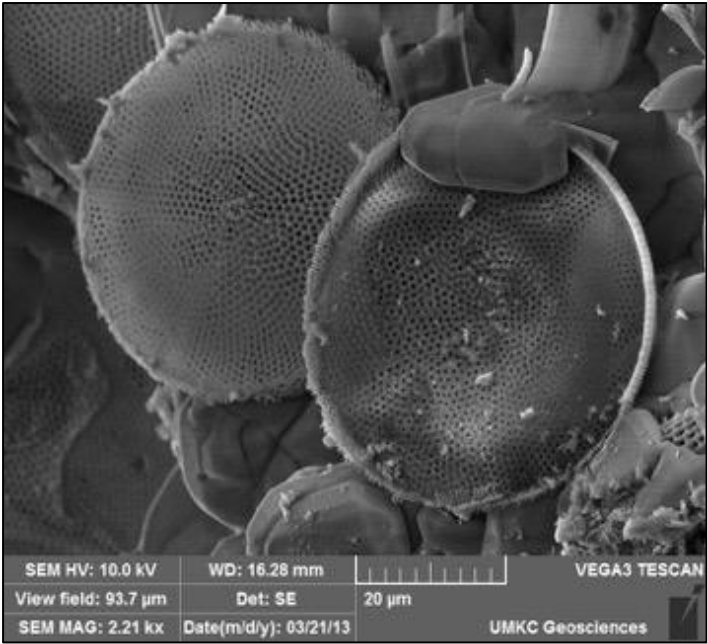


Aulacoseira lirata and *Stephanodiscus* sp.

Section 3 (20.05-20.14 m)



Aulacoseira lirata and *Stephanodiscus* sp.



Stephanodiscus sp.

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VITA

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